

“I GET ALL THE NEWS I NEED ON THE WEATHER REPORT”

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Francois Engelbrecht
Stephanie Landman

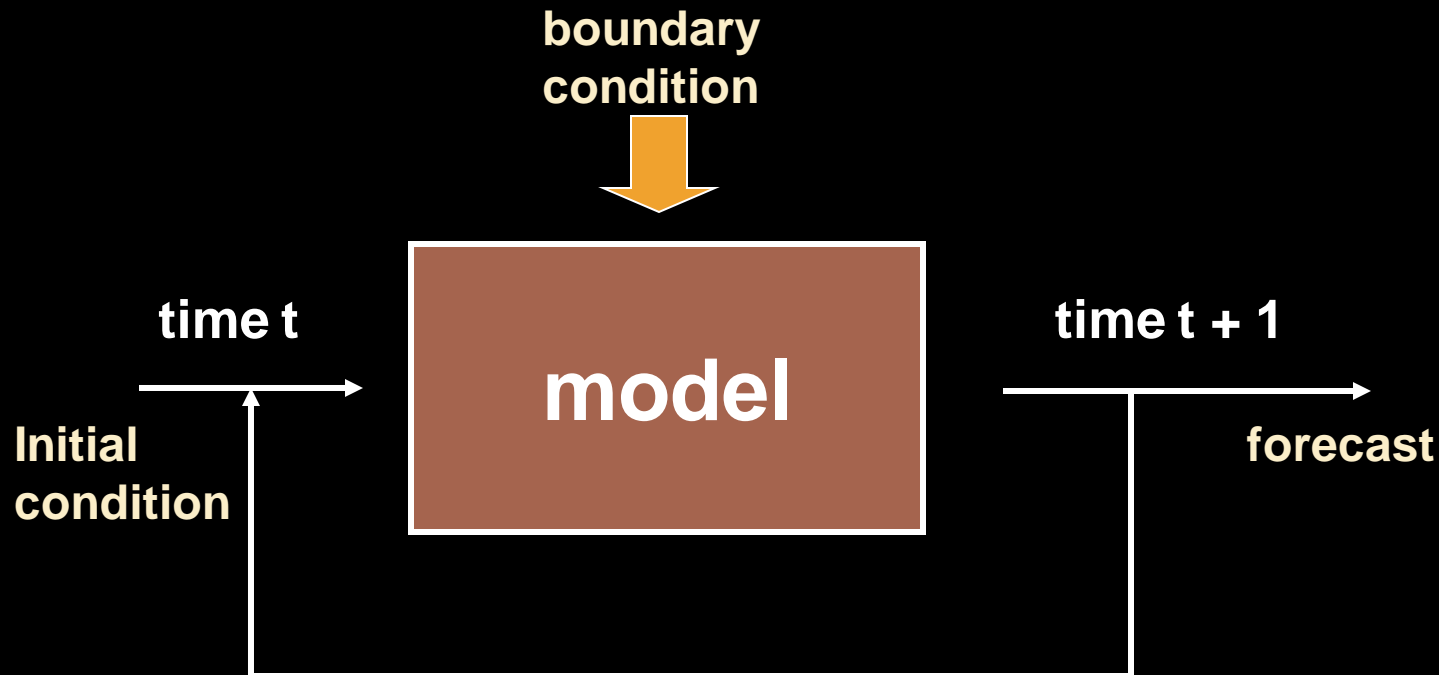


WMO FORECAST TIME RANGES

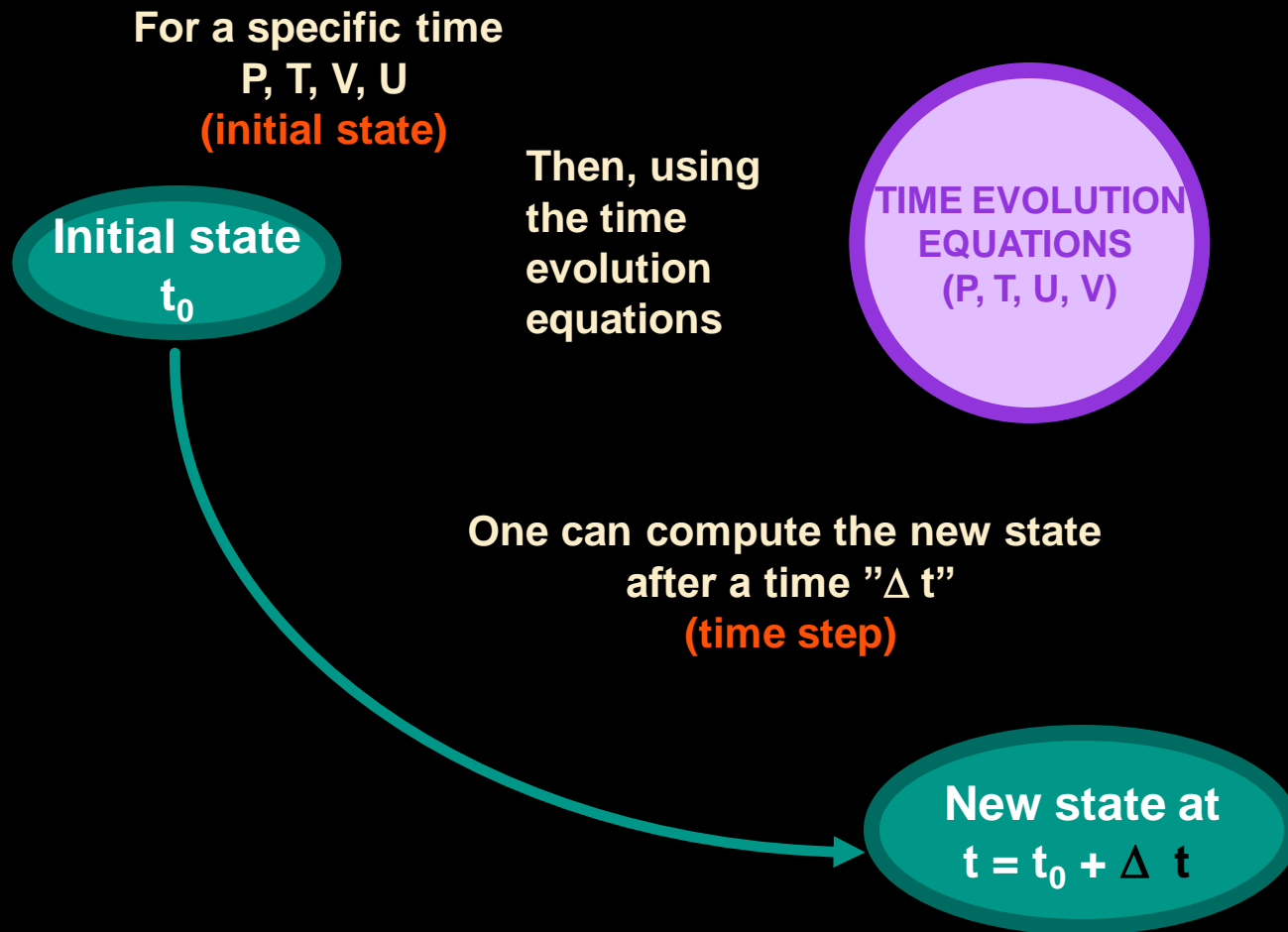
- × **Nowcasting:** A description of current weather parameters and 0 to 2 hours' description of forecast weather parameters
- × **Very short-range weather forecasting:** Up to 12 hours' description of weather parameters
- × **Short-range weather forecasting:** Beyond 12 hours' and up to 72 hours' description of weather parameters
- × **Medium-range weather forecasting:** Beyond 72 hours' and up to 240 hours' description of weather parameters
- × **Extended-range weather forecasting:** Beyond 10 days' and up to 30 days' description of weather parameters. Usually averaged and expressed as a departure from climate values for that period
- × **Long-range forecasting:** From 30 days up to two years
 - + **Month forecast:** Description of averaged weather parameters expressed as a departure (deviation, variation, anomaly) from climate values for that month at any lead-time
 - + **Seasonal forecast:** Description of averaged weather parameters expressed as a departure from climate values for that season at any lead-time
- × **Climate forecasting:** Beyond two years
 - + **Climate variability prediction:** Description of the expected climate parameters associated with the variation of interannual, decadal and multi-decadal climate anomalies
 - + **Climate prediction:** Description of expected future climate including the effects of both natural and human influences

A NUMERICAL MODEL: OPERATIONAL ORGANIZATION (1)

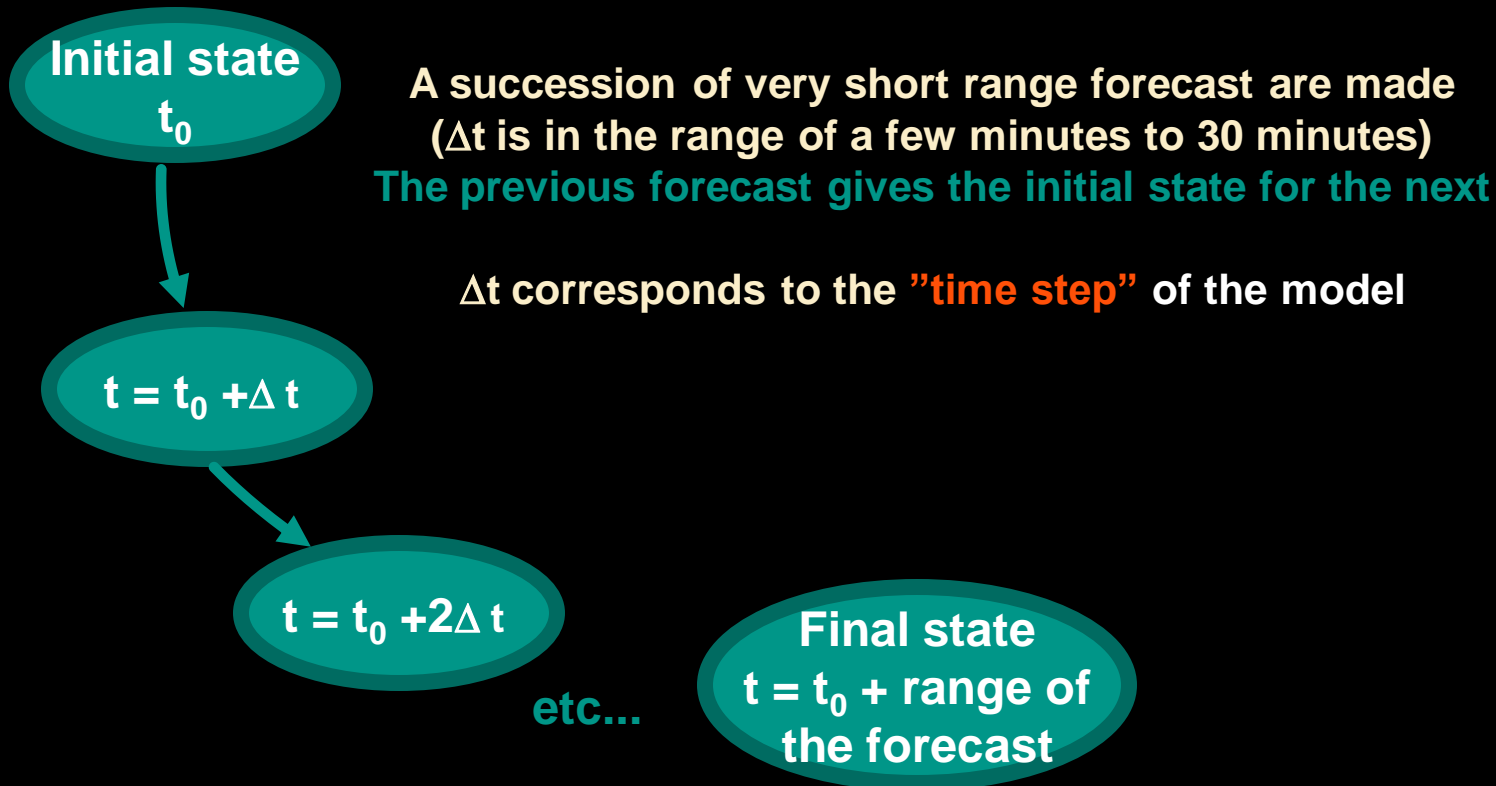
A typical run:



A NUMERICAL MODEL: OPERATIONAL ORGANIZATION (2)



A NUMERICAL MODEL: OPERATIONAL ORGANIZATION (3)



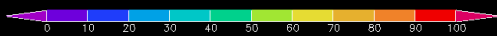
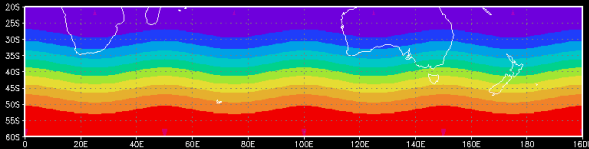
NOT SO LONG AGO...

1980s

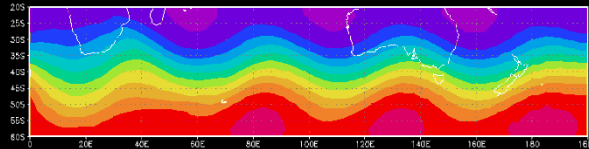
1940s

$$\frac{\partial}{\partial t} \nabla^2 \psi = -\mathbf{V}_\psi \cdot \nabla (\nabla^2 \psi + f)$$

Streamfunction Field at 0 Hours



Streamfunction Field at 120 Hours



Sonale momentumvergelyking

$$\frac{U}{a \cos \phi} \frac{\partial u^*}{\partial \lambda} + \frac{v^*}{a} \frac{\partial U}{\partial \phi} + \omega^* \frac{\partial U}{\partial p} - f v^* + \frac{1}{a \cos \phi} \frac{\partial \Phi^*}{\partial \lambda} - F_{(x)}^* = 0$$

Meridionale momentumvergelyking

$$\frac{U}{a \cos \phi} \frac{\partial v^*}{\partial \lambda} + f u^* + \frac{1}{a} \frac{\partial \Phi^*}{\partial \phi} - F_{(y)}^* = 0$$

Kontinuiteitsvergelyking

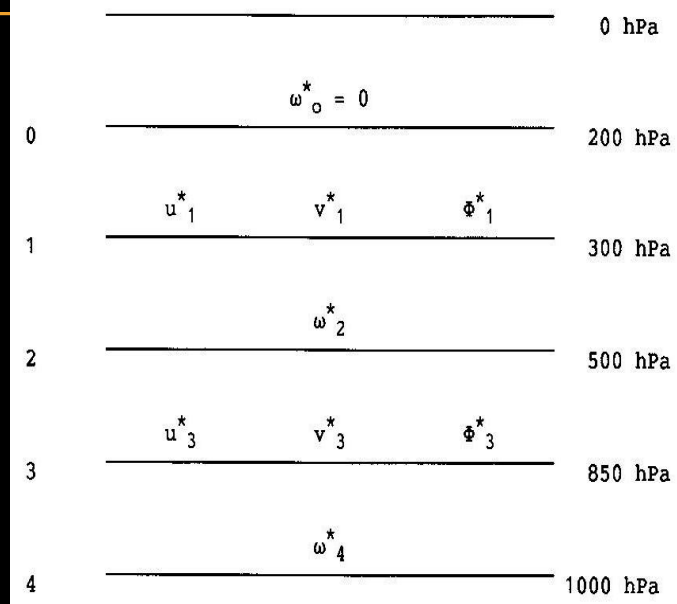
$$\frac{1}{a \cos \phi} \frac{\partial u^*}{\partial \lambda} + \frac{1}{a \cos \phi} \frac{\partial (v^* \cos \phi)}{\partial \phi} + \frac{\partial \omega^*}{\partial p} = 0$$

Hidrostatiese vergelyking

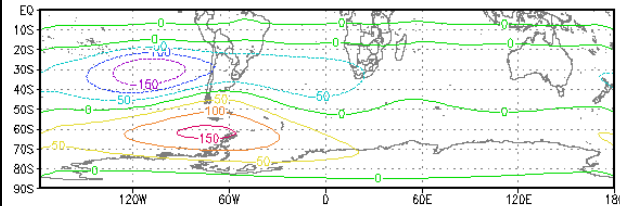
$$\frac{\partial \Phi^*}{\partial p} + \frac{RT^*}{p} = 0$$

Termodinamiese vergelyking

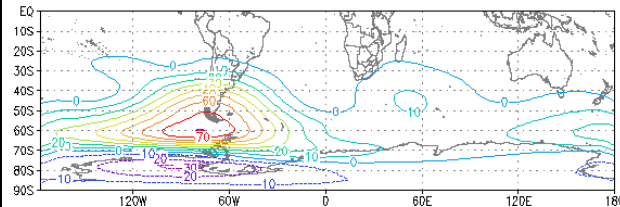
$$\frac{U}{a \cos \phi} \frac{\partial T^*}{\partial \lambda} + \frac{v^*}{a} \frac{\partial [T]}{\partial \phi} - S_p \omega^* = \frac{Q^*}{c_p}$$



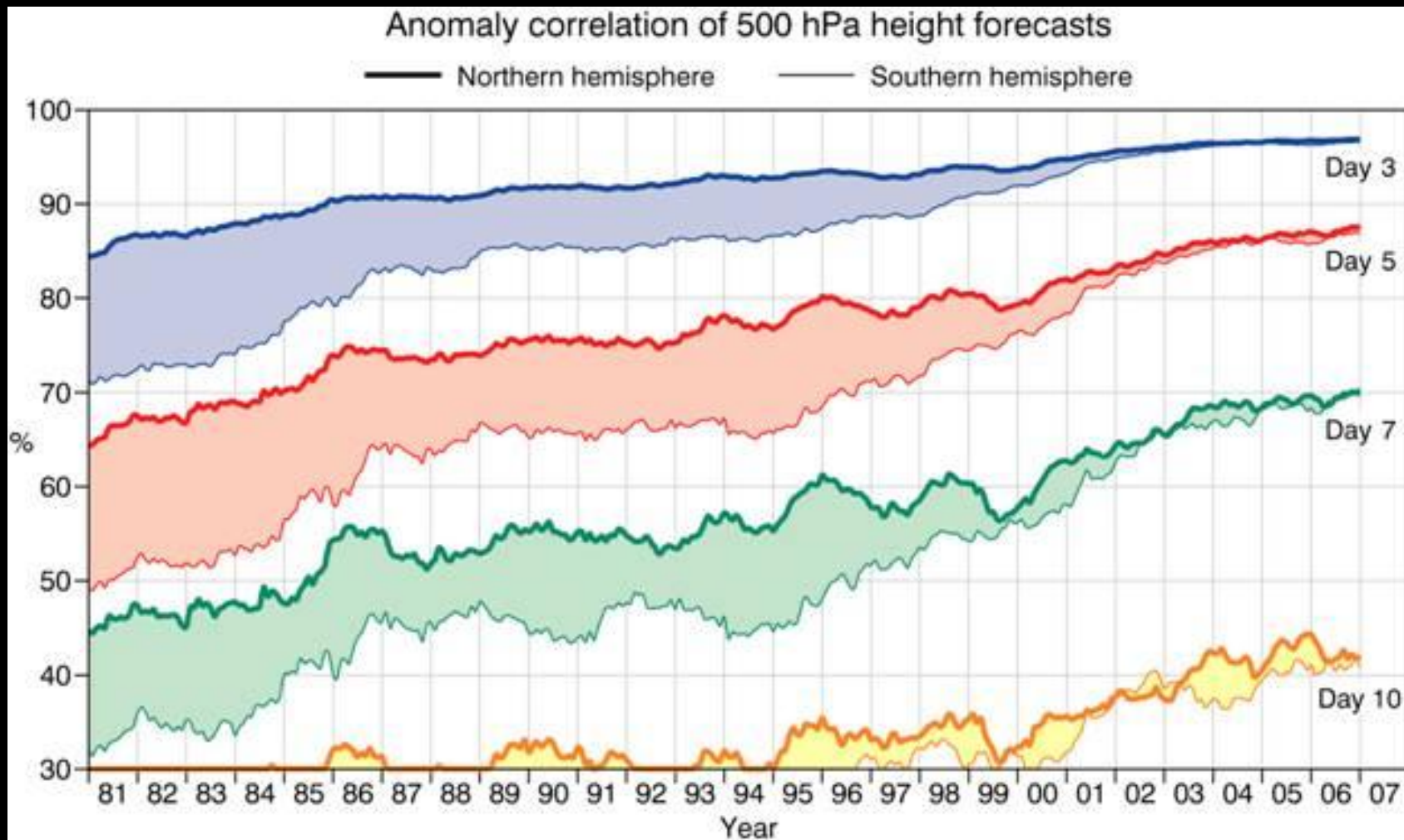
EDDY GEOPOTENTIAL HEIGHT RESPONSE TO DIABATIC HEATING: NINO3.4
300 hPa



850 hPa

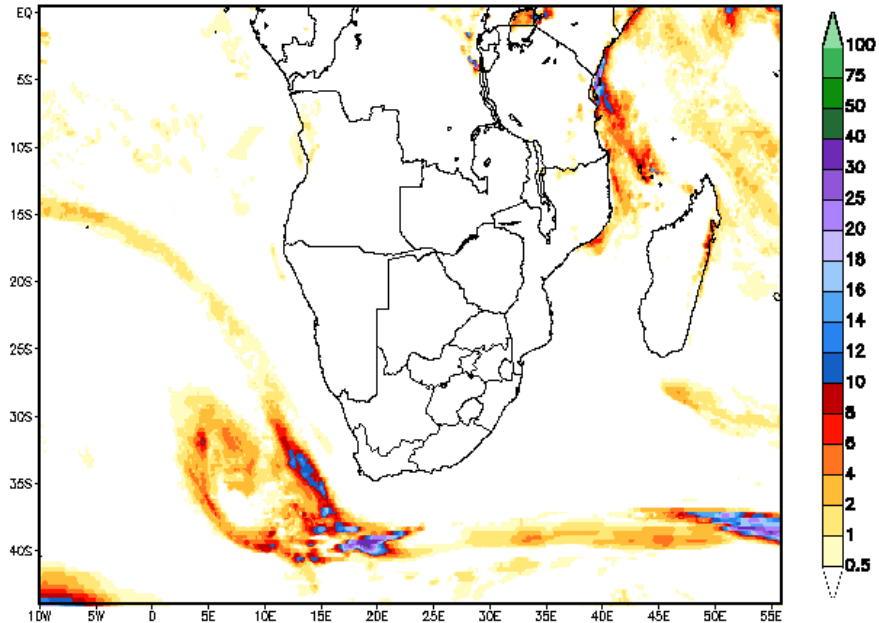


CONVERGENCE OF NWP SKILL



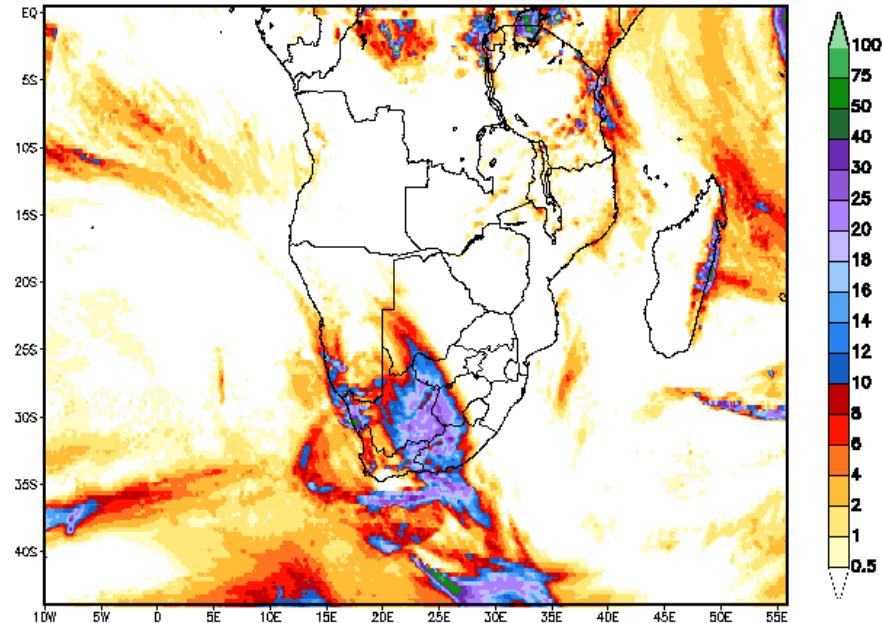
SHORT-RANGE WEATHER FORECAST MODELS

UM 12km horizontal resolution – xaana Run:
Accumulating Precipitation for past 6 hours (mm)



Total precipitation of 00Z to 06Z, 10 JUL – Initiated 00Z 10 JUL 2009

UM 12km horizontal resolution – xaana Run:
Accumulating Precipitation for past 24 hours (mm)



Total precipitation of 00Z to 24Z, 11 JUL – Initiated 00Z 10 JUL 2009

EFFECTS OF SMALL ERRORS IN THE INITIAL CONDITIONS

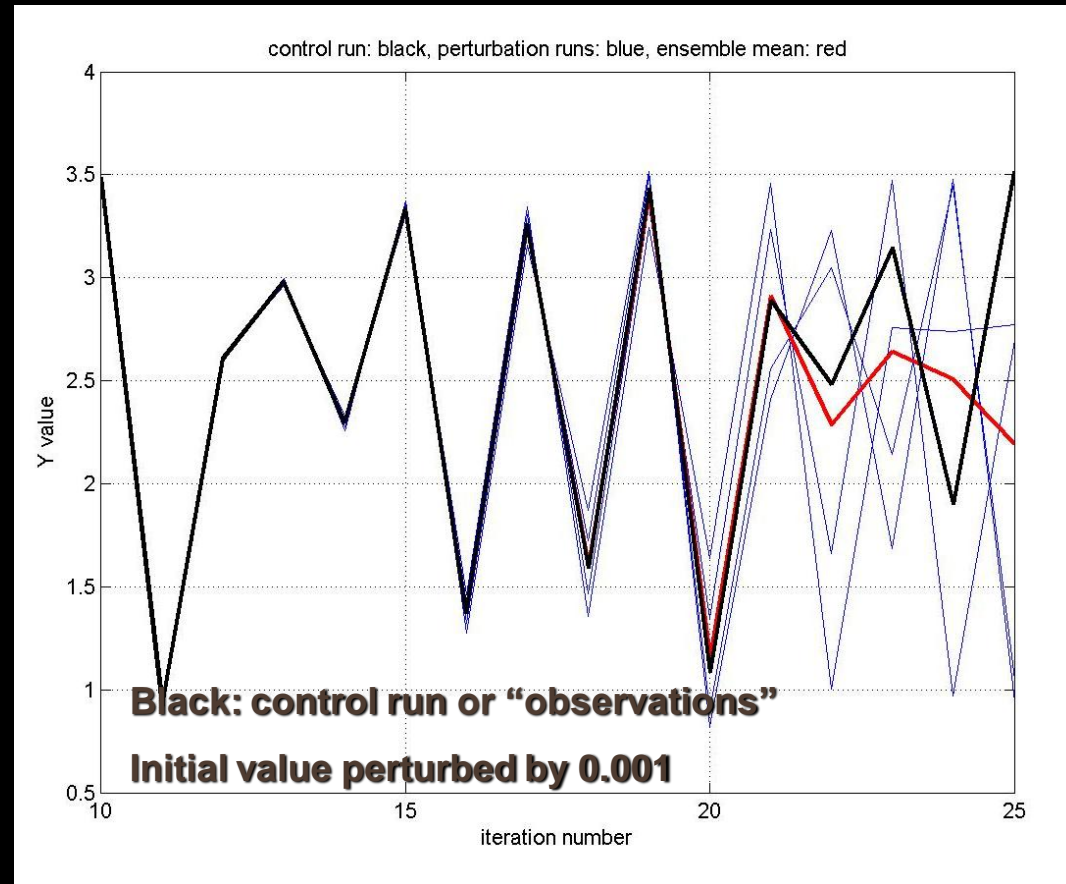
- ✘ There will always be atmospheric motions whose scales are too small to be properly observed and represented in a model
- ✘ Thus, there is an unavoidable level of error in the determination of the initial state
- ✘ The nonlinearity and instability of the atmospheric flow will inevitably cause the small inherent errors in the initial data to grow and gradually affect the larger scales of motion so that the forecast flow field will eventually evolve differently from the actual flow

FIRST-ORDER QUADRATIC DIFFERENCE EQUATION

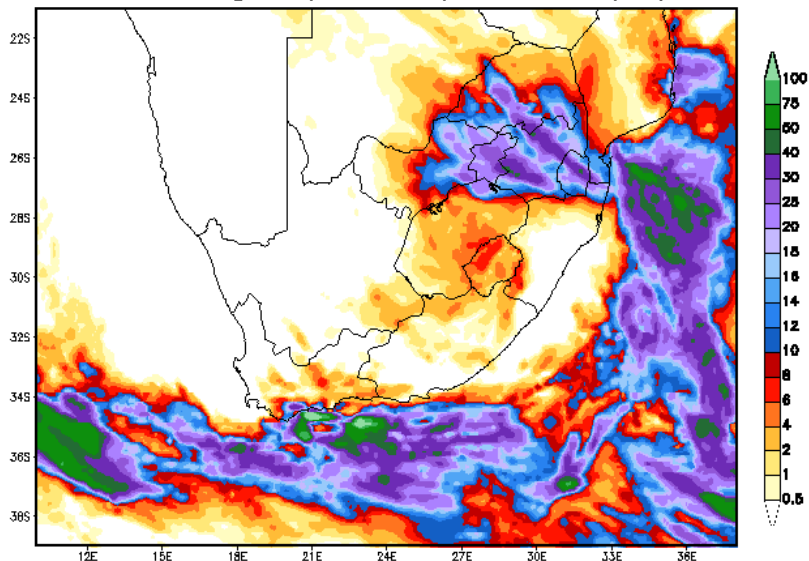
- ✘ Lorenz illustrated the general problem of predictability by considering the first-order quadratic difference equations:

$$Y_{s+1} = aY_s - Y_s^2$$

- ✘ Figure is for $Y(0) = 1.5$;
 $a = 3.75$

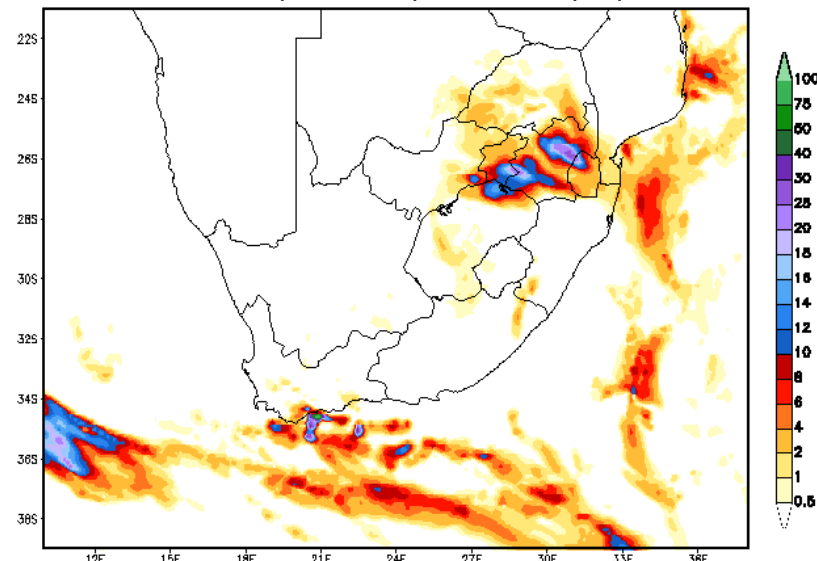


UM 12km horizontal resolution – xaang Run:
Accumulating Precipitation for past 24 hours (mm)



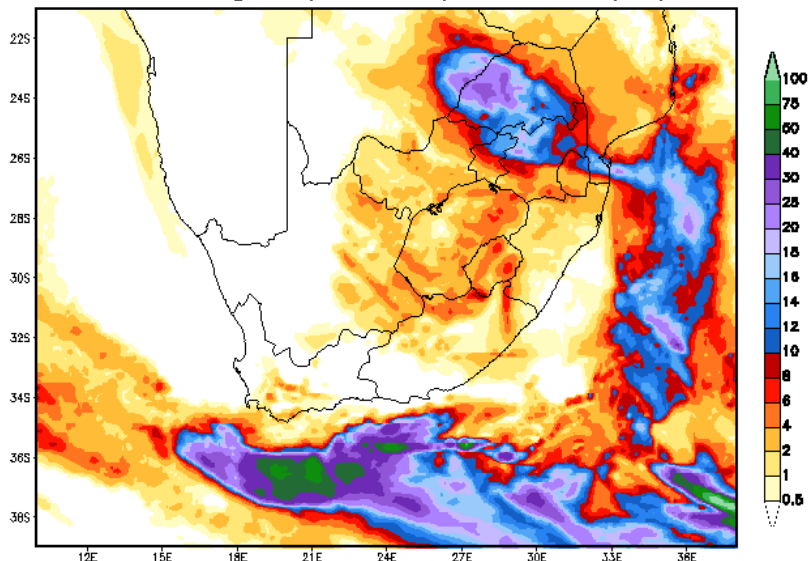
Total precipitation of 00Z to 24Z, 11 JUN – Initiated 00Z 11 JUN 2009

UM 12km horizontal resolution – xaang Run:
Total Precipitation for past 3 hours (mm)



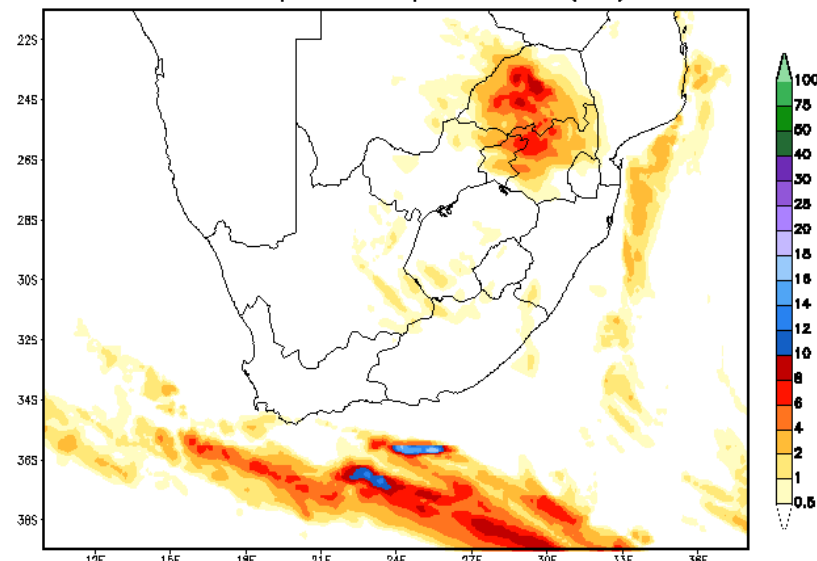
Total precipitation of 06Z to 09Z, 11 JUN – Initiated 00Z 11 JUN 2009

UM 12km horizontal resolution – xaana Run:
Accumulating Precipitation for past 24 hours (mm)



Total precipitation of 00Z to 24Z, 11 JUN – Initiated 00Z 11 JUN 2009

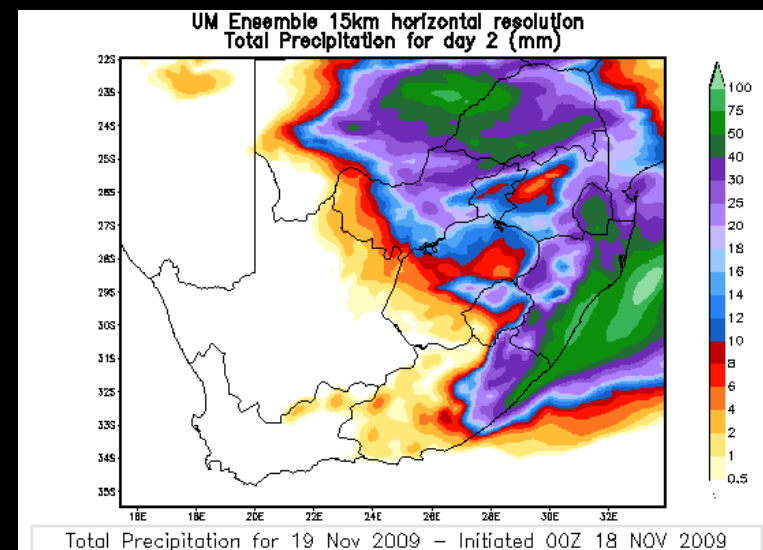
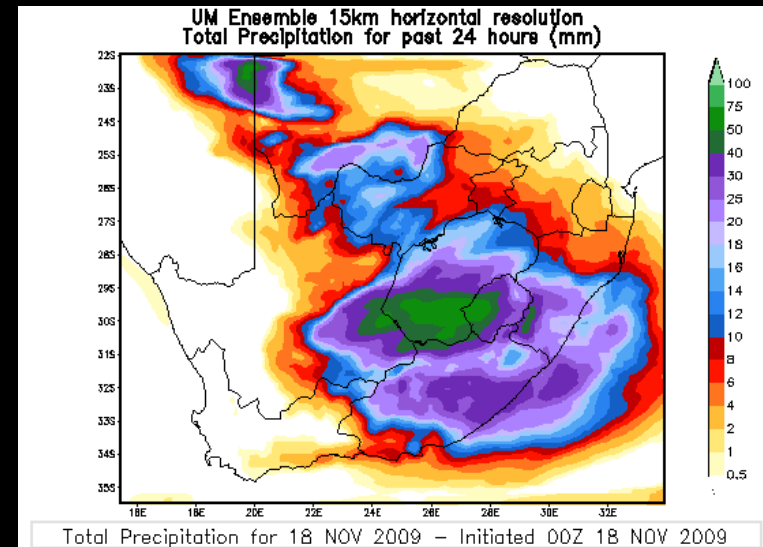
UM 12km horizontal resolution – xaana Run:
Total Precipitation for past 3 hours (mm)



Total precipitation of 06Z to 09Z, 11 JUN – Initiated 00Z 11 JUN 2009

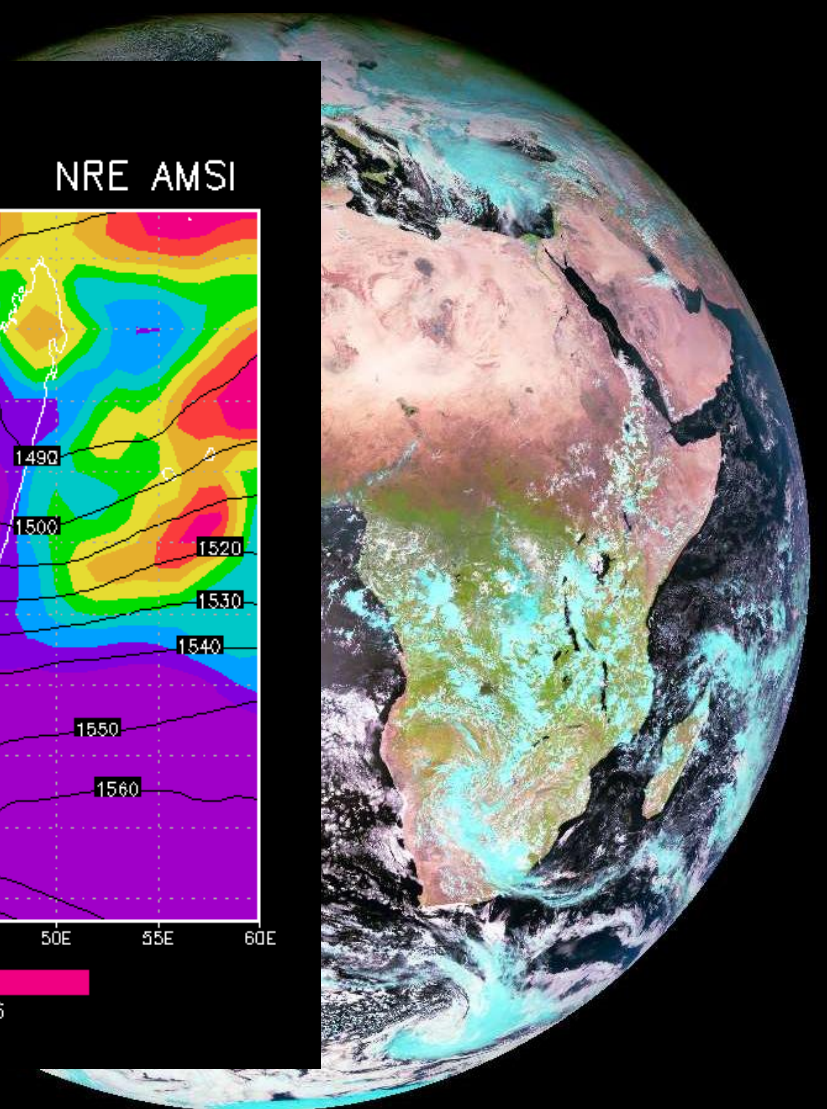
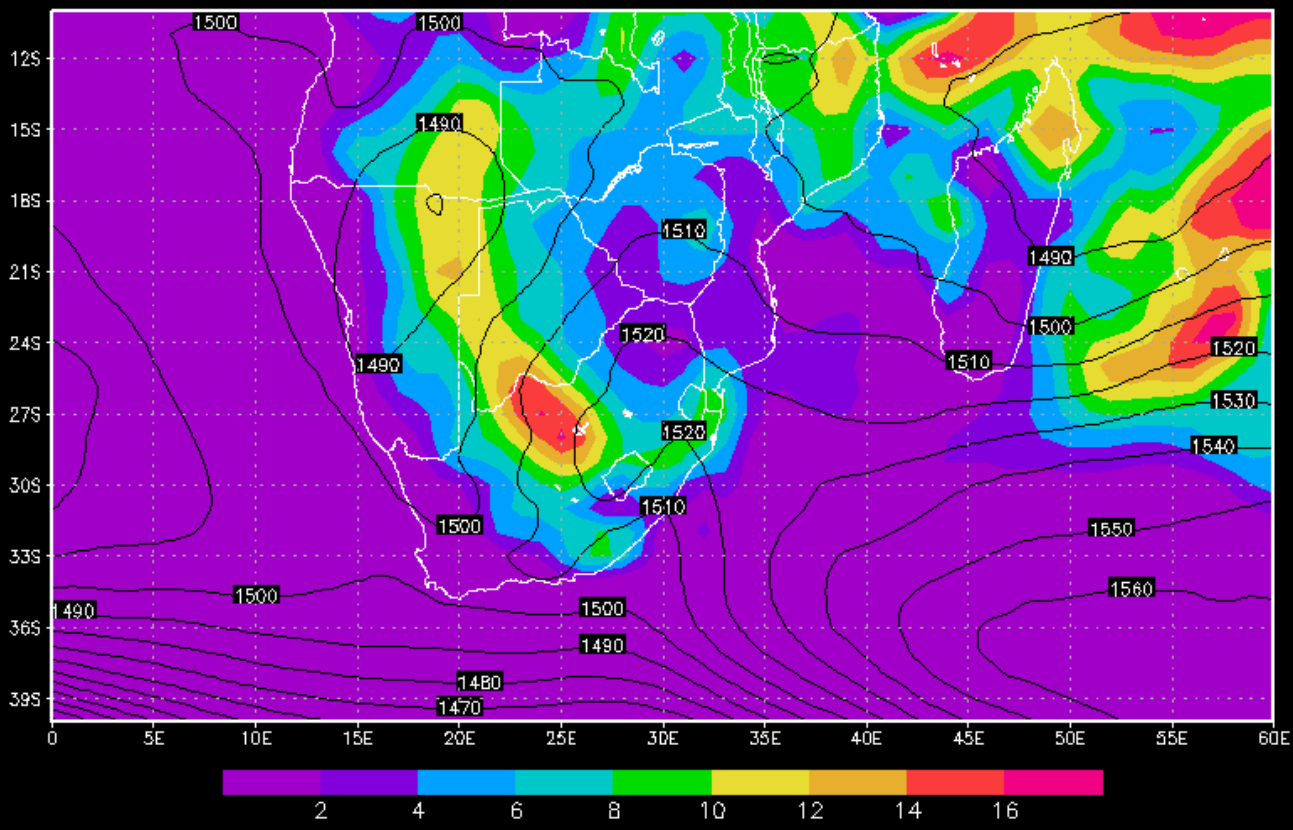
NEW MULTI-MODEL SHORT-RANGE ENSEMBLE SYSTEM (PRECIPITATION)

- ✘ 24hour totals for:
 - + day1 (14 members) and
 - + day2 (6 members)
- ✘ Unified Model (different configurations and resolutions)
 - + 10 members:
 - ✘ 12km (xaana/ng/nj)
 - ✘ 15km (xaaha/hc)
- ✘ WRF model
 - + 2 members:
 - ✘ 12km
 - ✘ Non-hydrostatic mesoscale core
 - + 2 members:
 - ✘ 15km
 - ✘ Advanced Research WRF core



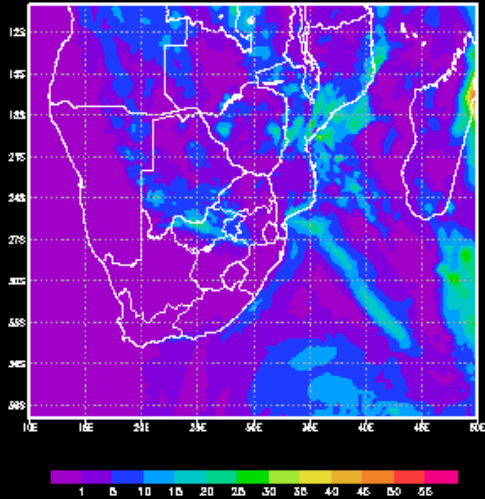
RAINFALL FORECAST FOR 22 JAN 10

CCAM Rainfall forecast (mm) for 2010-01-22 NRE AMSI

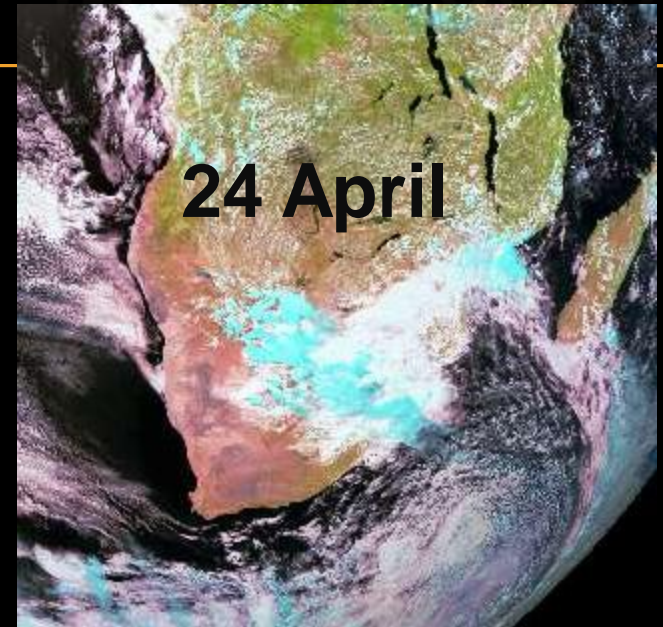
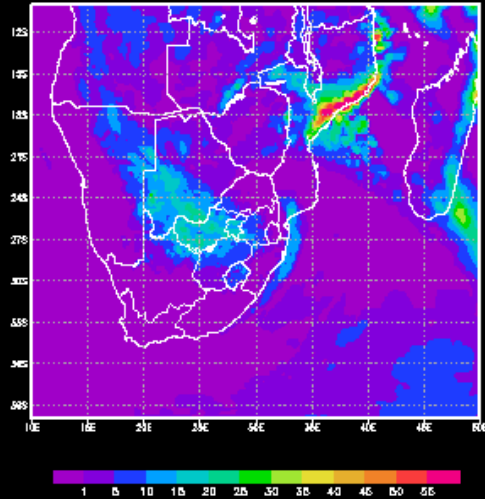


RECENT COLD/WET SPELL IN APRIL 2010

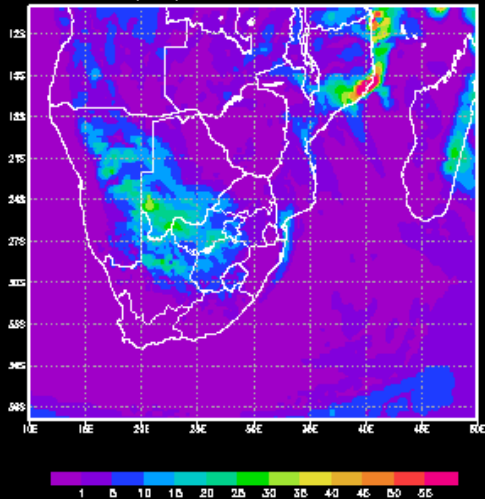
CCAM md (mm) 2010-04-23 CSIR-NRE AMSI



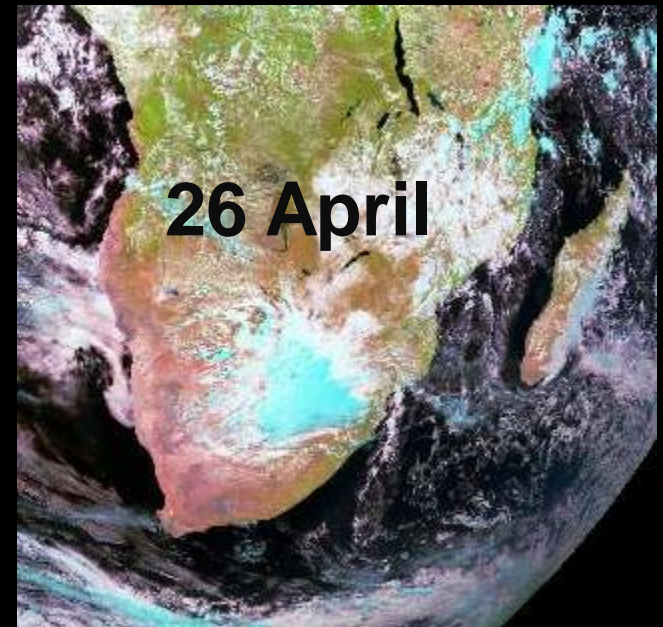
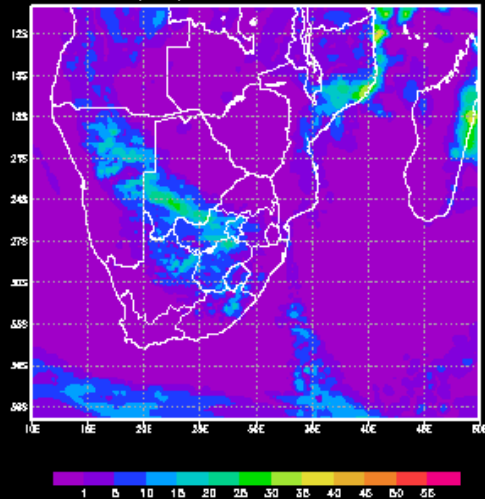
CCAM md (mm) 2010-04-24 CSIR-NRE AMSI

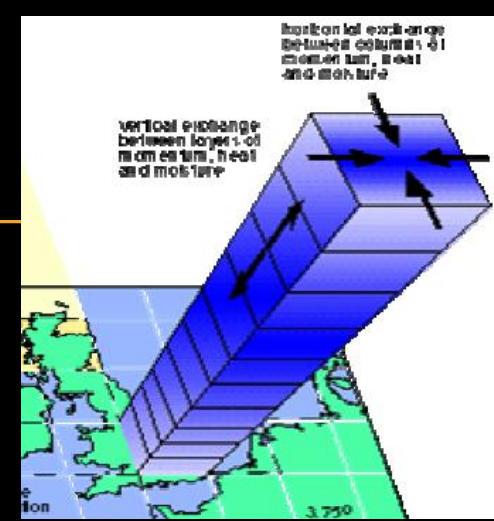
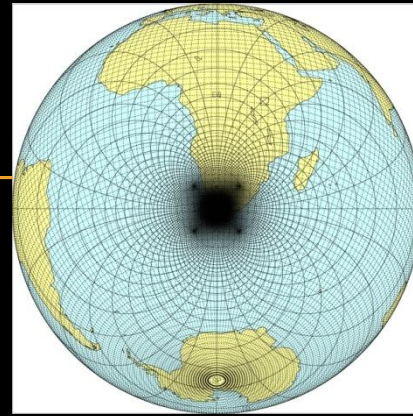
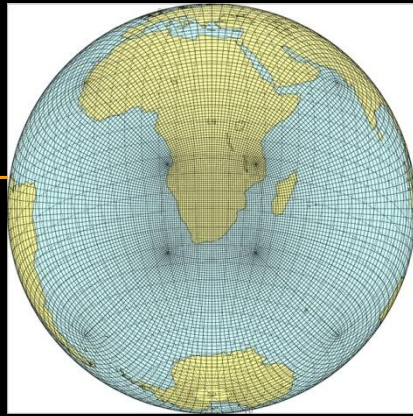
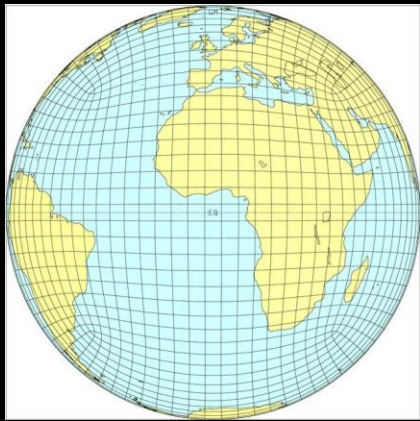


CCAM md (mm) 2010-04-25 CSIR-NRE AMSI



CCAM md (mm) 2010-04-26 CSIR-NRE AMSI



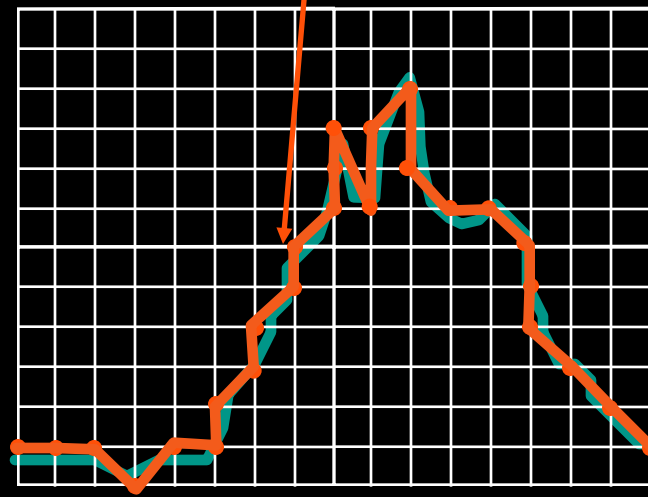
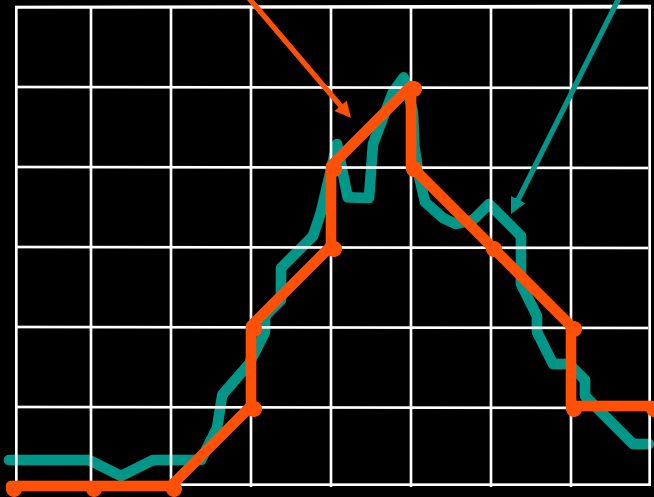


An example for the **topography** of the model

Real topography profile

Topography profile in the model

Topography profile in the model using a doubling of horizontal and vertical resolution

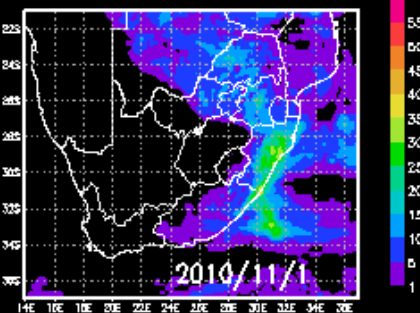
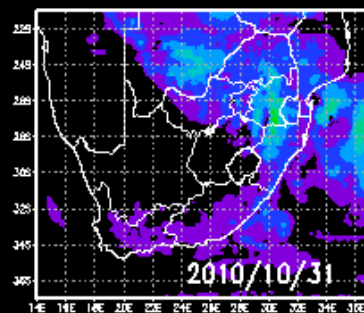
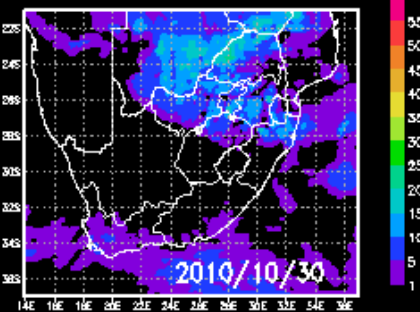
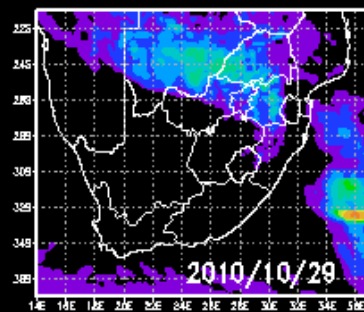
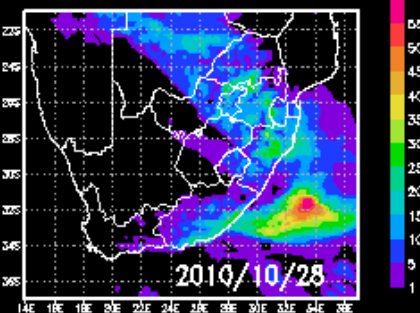
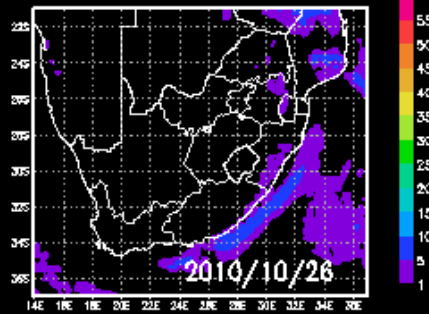
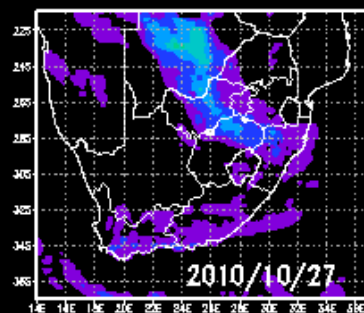


Latest weather forecast

CCAM rainfall (mm)

Produced by:

The Climate Studies and Modelling
Group (CSM) of the CSIR.

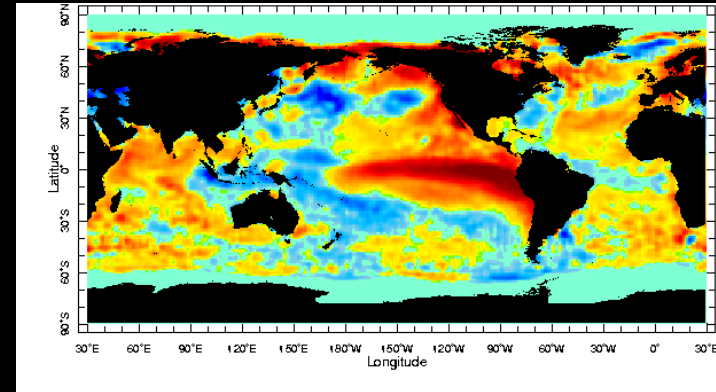


LIMITS OF LONGER RANGE FORECASTS

- ✘ Great progress has been made to predict the day-to-day state of the atmosphere (e.g., frontal movement, winds, pressure)
- ✘ However, day-to-day fluctuations in weather are not predictable beyond two weeks
- ✘ Beyond that time, errors in the data defining the state of the atmosphere at the start of a forecast period grow and overwhelm valid forecast information
- ✘ This so called “**chaotic**” behaviour is an *inherent property* of the atmosphere

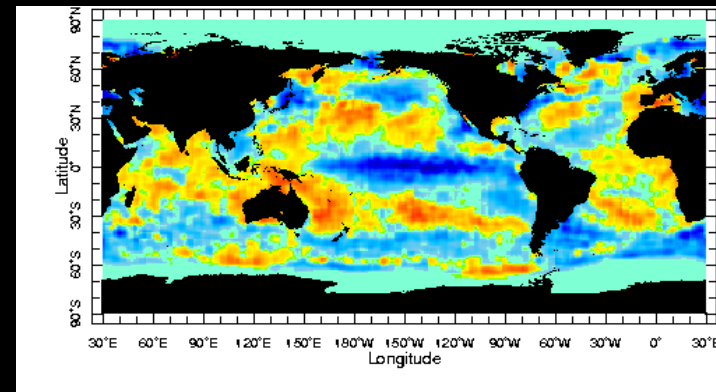
HOW IS IT THEN POSSIBLE TO PREDICT **SEASONAL** CLIMATE ANOMALIES?

Predictions of rainfall, frontal passages, etc. for a particular day at a certain location several months ahead has no usable skill. However, there is some skill in predicting anomalies in the **seasonal average** of the weather. The predictability of seasonal climate anomalies results primarily from the influence of **slowly evolving boundary conditions**, and most notably SSTs (i.e., El Niño and La Niña), on the atmospheric circulation.



Sea-surface temperature (SST) anomalies of September 1997 (El Niño of 1997/98)

Anomaly: departure from the mean or average



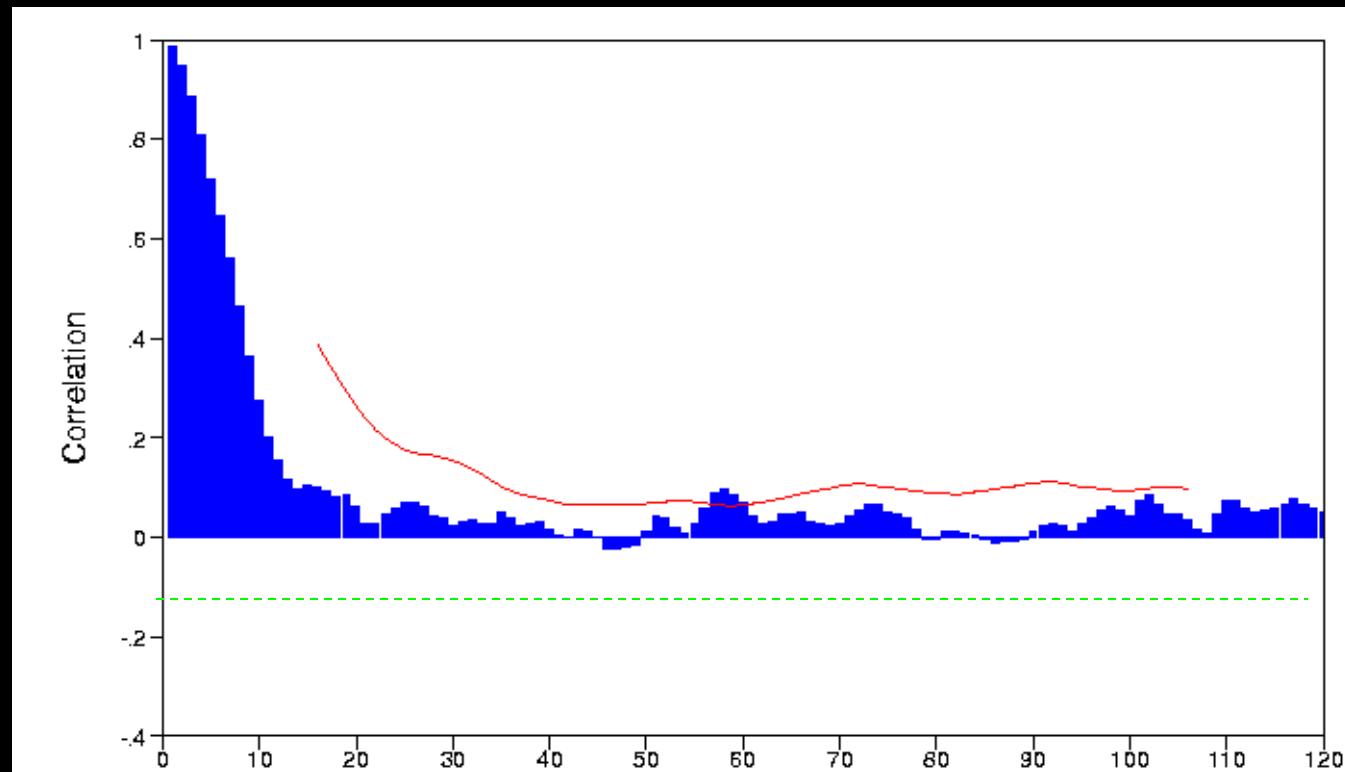
Sea-surface temperature (SST) anomalies of November 1988 (La Niña of 1988/89)

DYNAMICAL FORECASTS: SEASONAL FORECASTS

Daily Scores

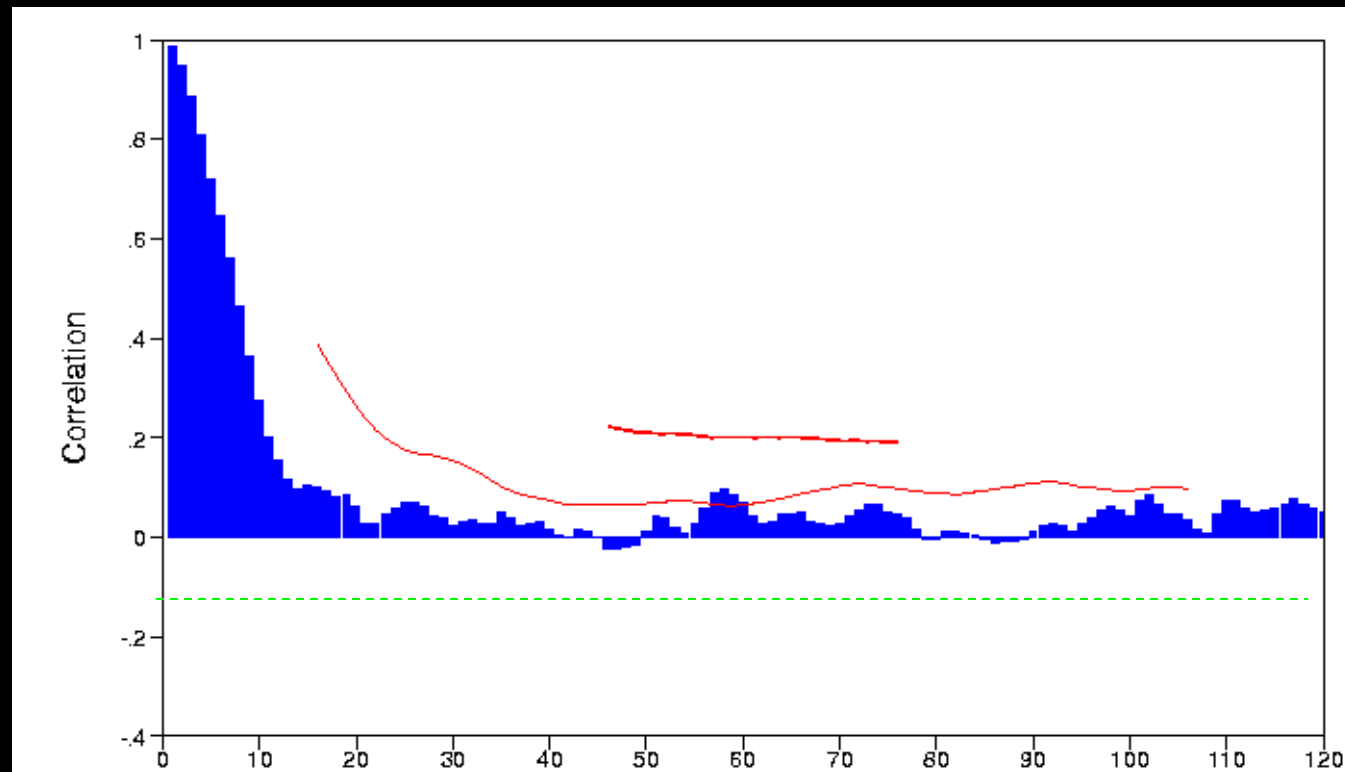
+

Monthly
running mean
Scores



DYNAMICAL FORECASTS: SEASONAL FORECASTS

Daily Scores
+
Seasonal
running mean
Scores

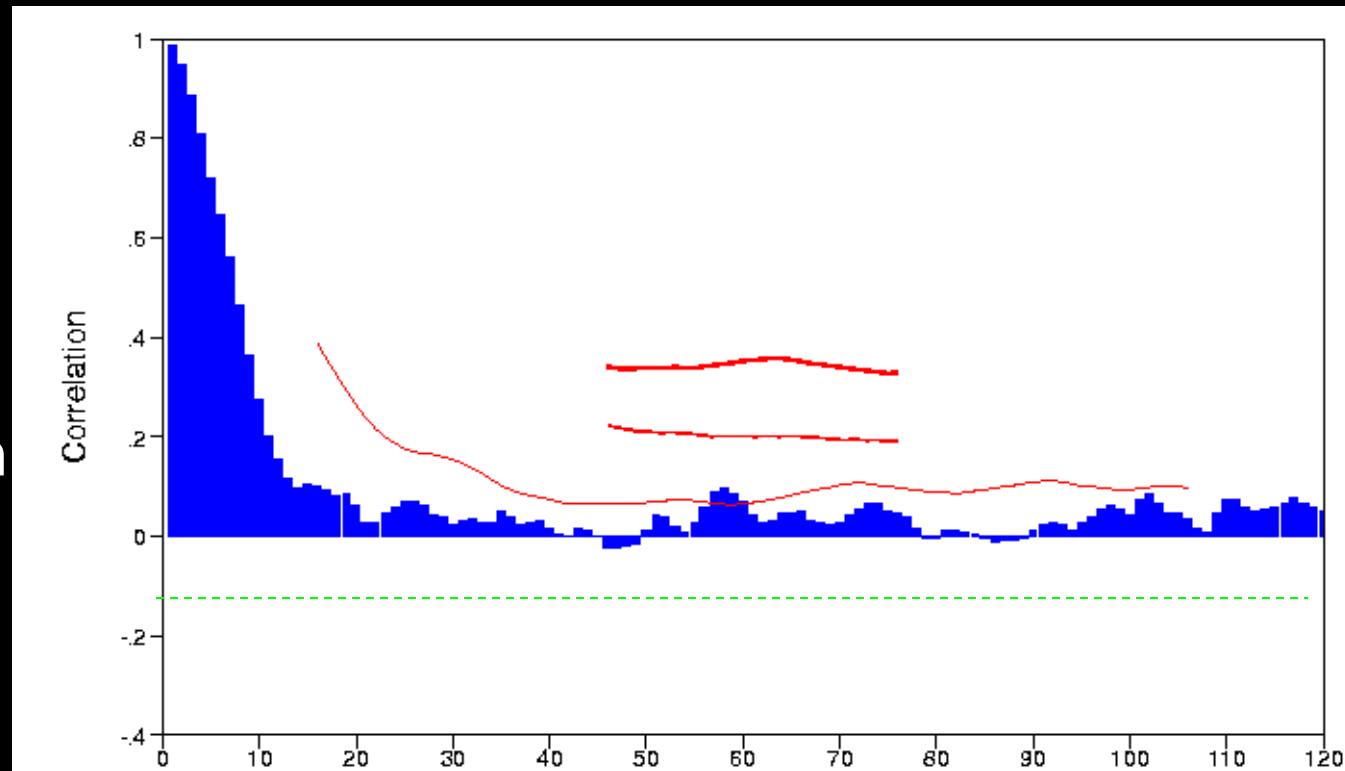


DYNAMICAL FORECASTS: SEASONAL FORECASTS

Daily Scores

+

Ensemble
forecast,
Seasonal
running mean
and SST
forecast



DYNAMICAL FORECASTS: SEASONAL FORECASTS

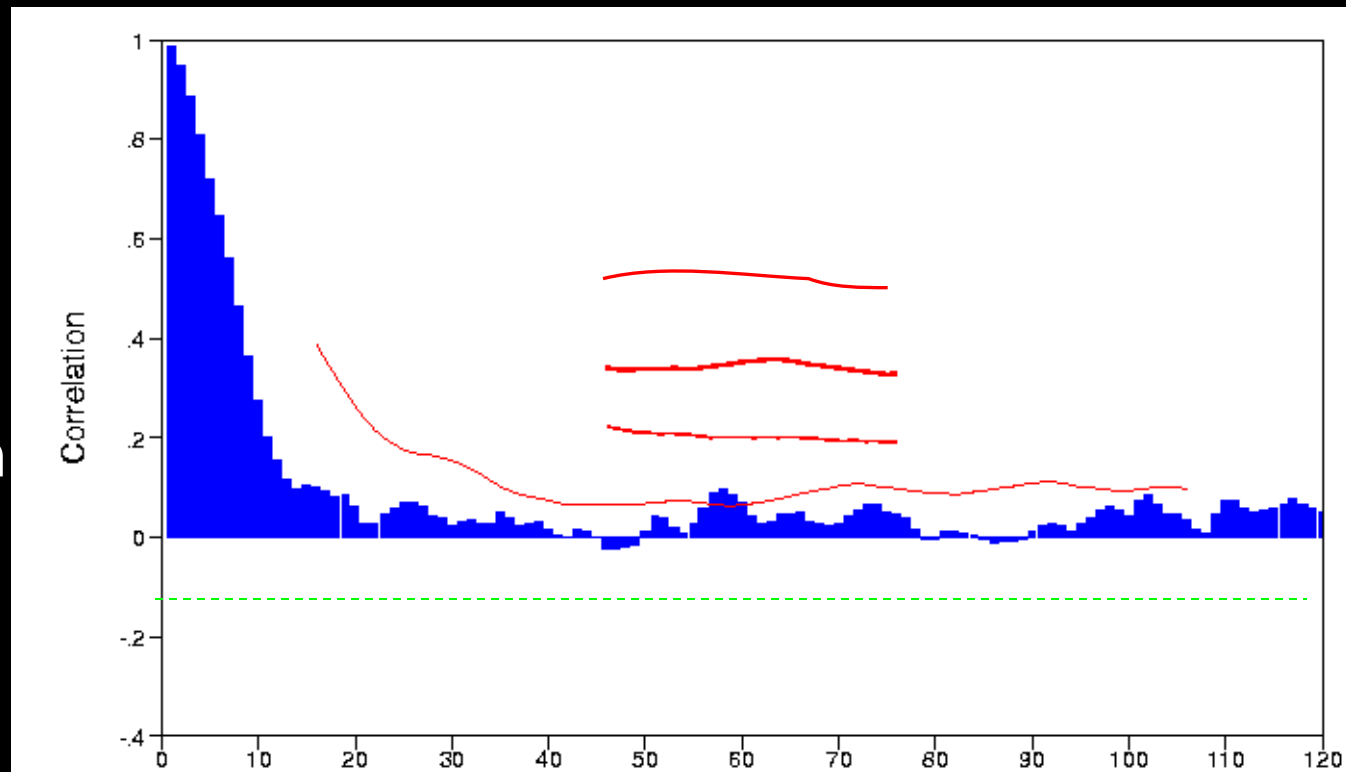
Daily Scores

+

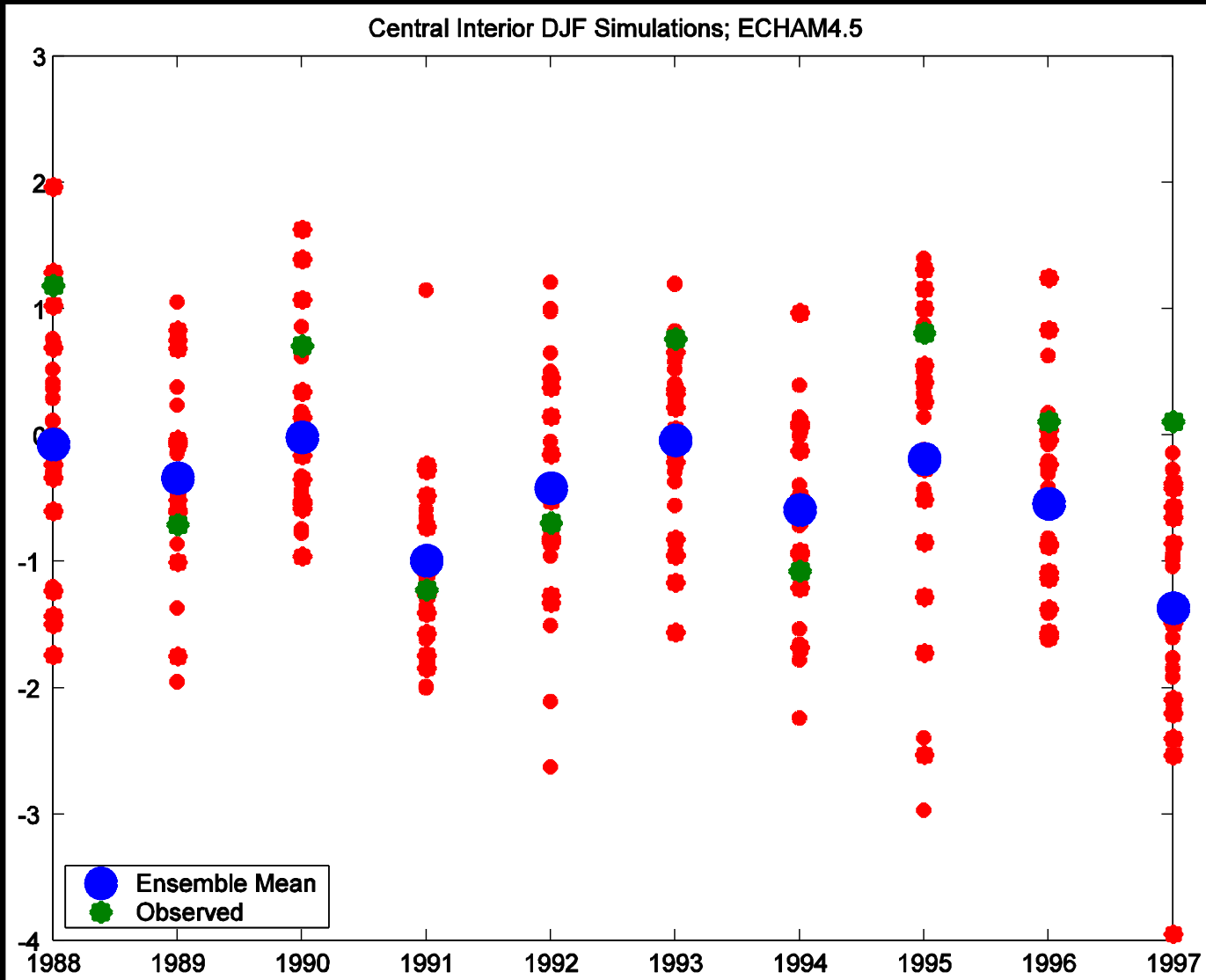
Ensemble
forecast,
Seasonal
running mean
and SST
forecast

+

MOS



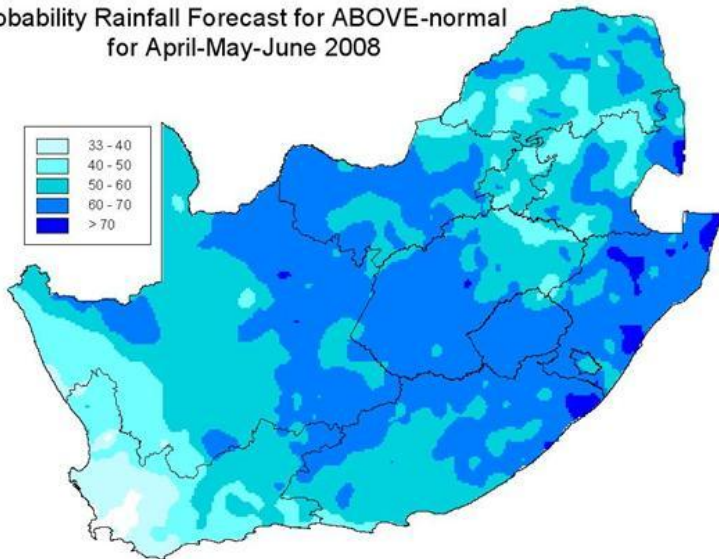
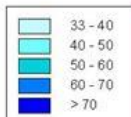
ENSEMBLE OF FORECASTS



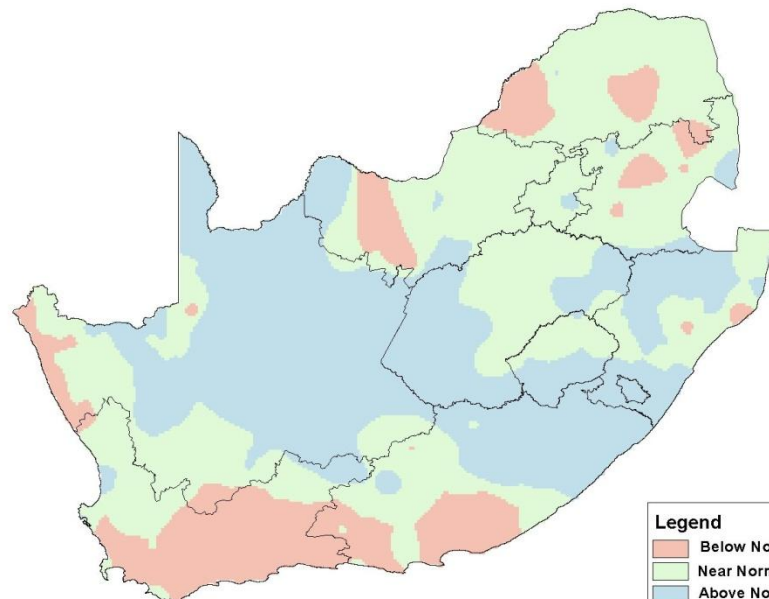
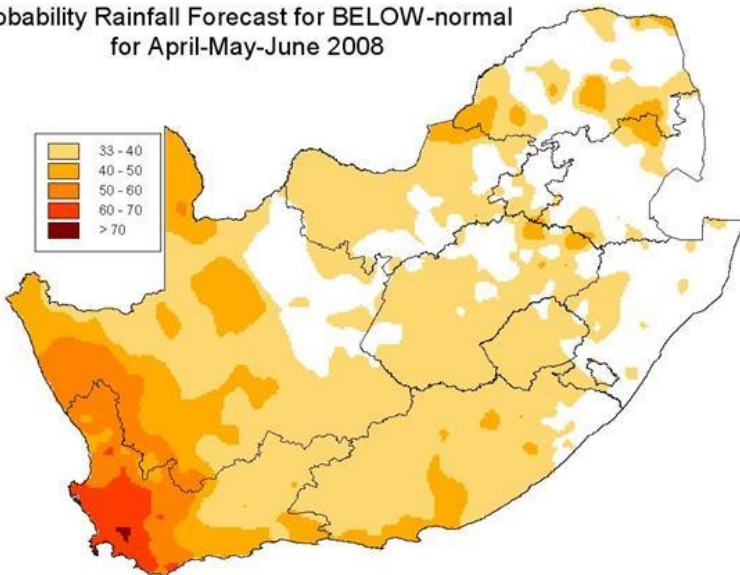
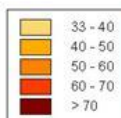
Objective multi-model forecast

Assessment of Rainfall for April to June 2008

Probability Rainfall Forecast for ABOVE-normal for April-May-June 2008

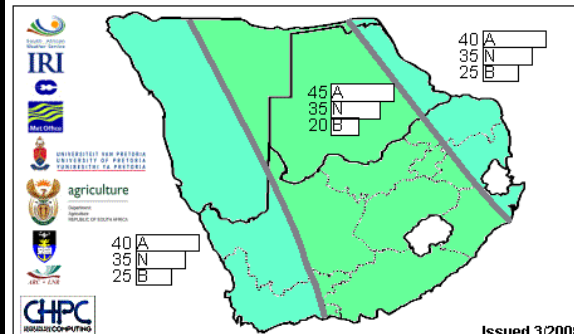


Probability Rainfall Forecast for BELOW-normal for April-May-June 2008

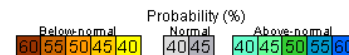


Subjective consensus forecast

Expected Total Rainfall for the period April-May-June 2008

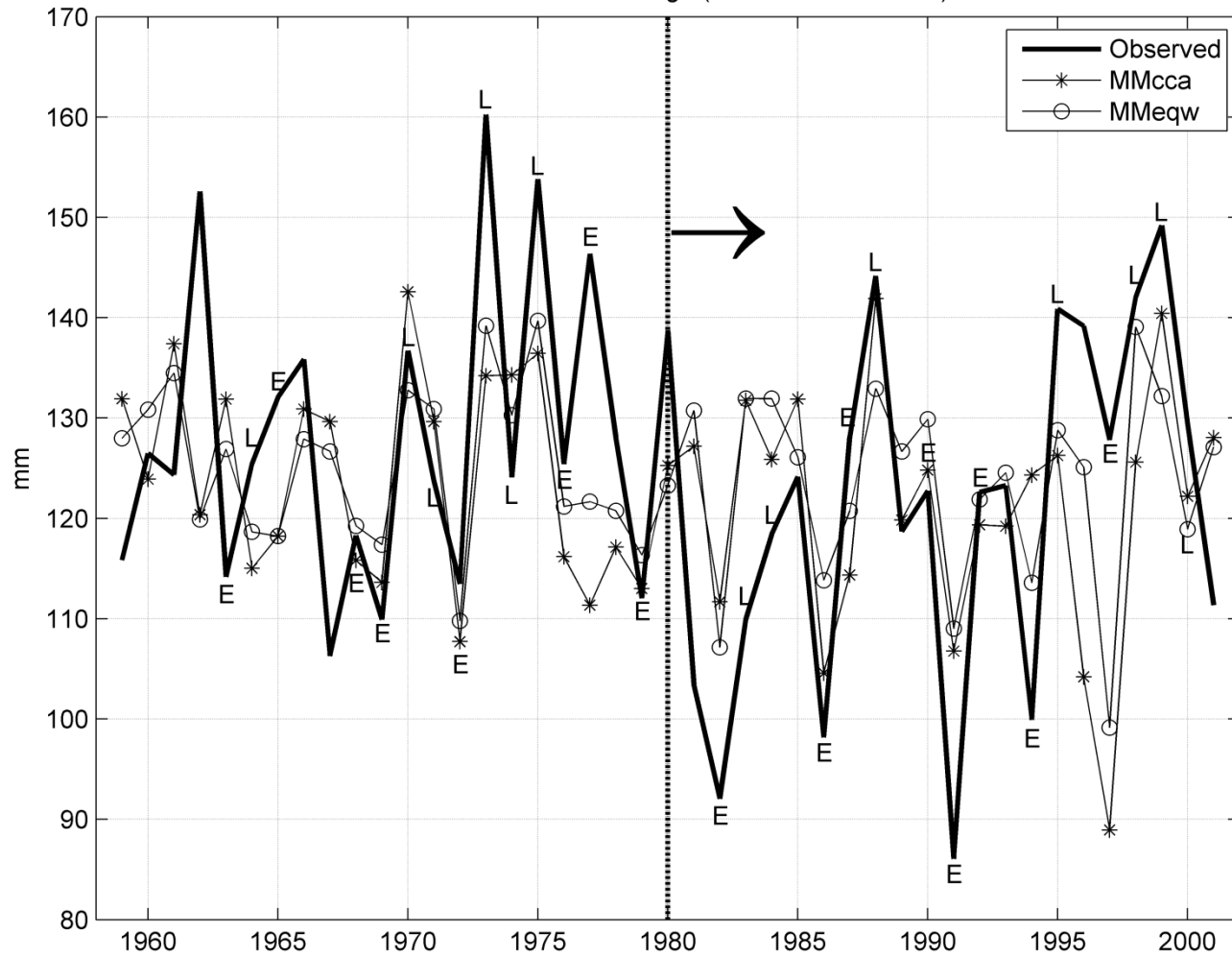


Issued 3/2008



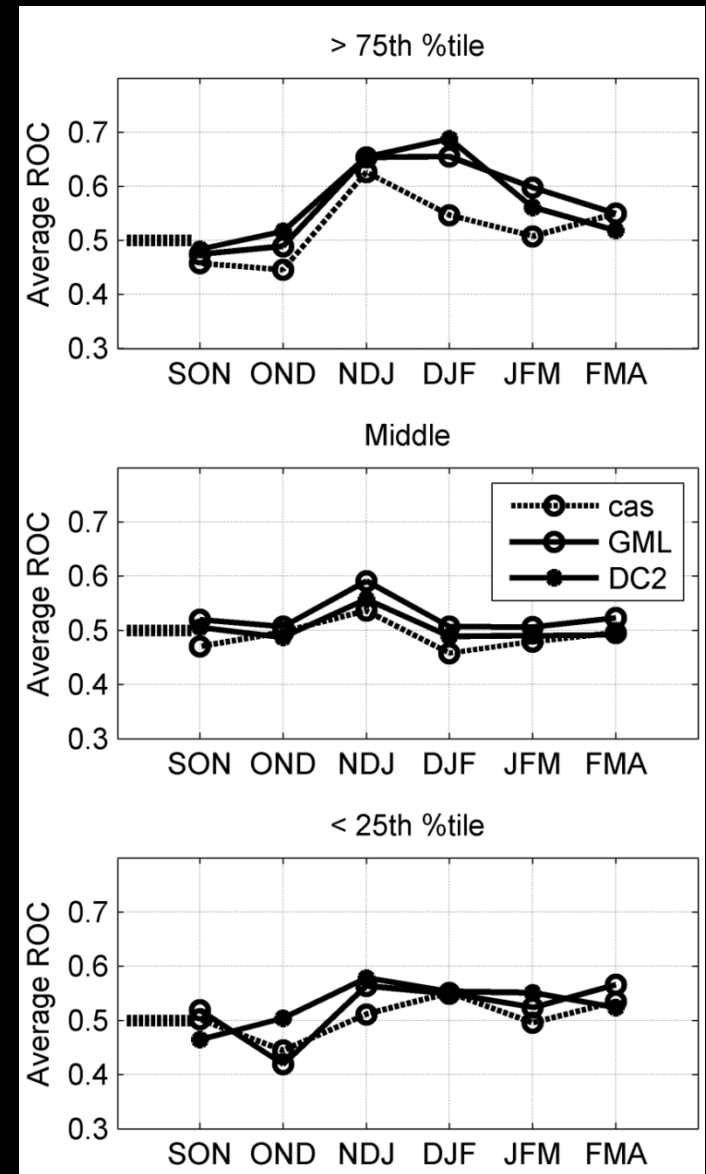
Please send comments to longrange@weathersa.co.za

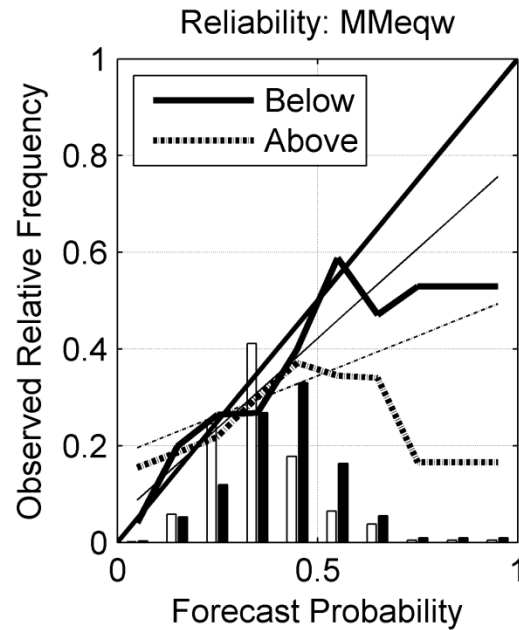
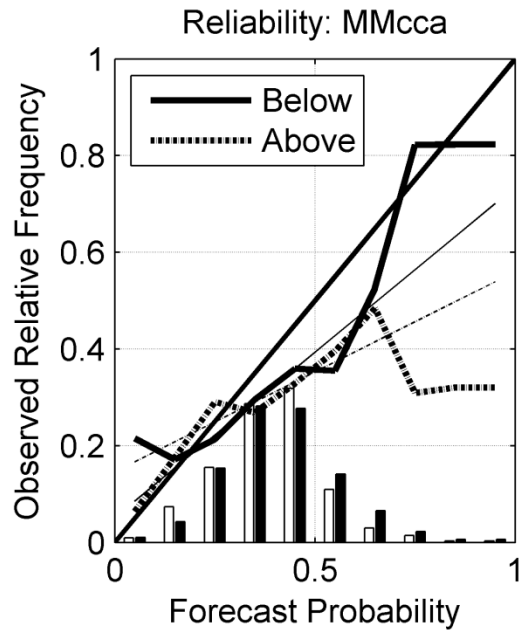
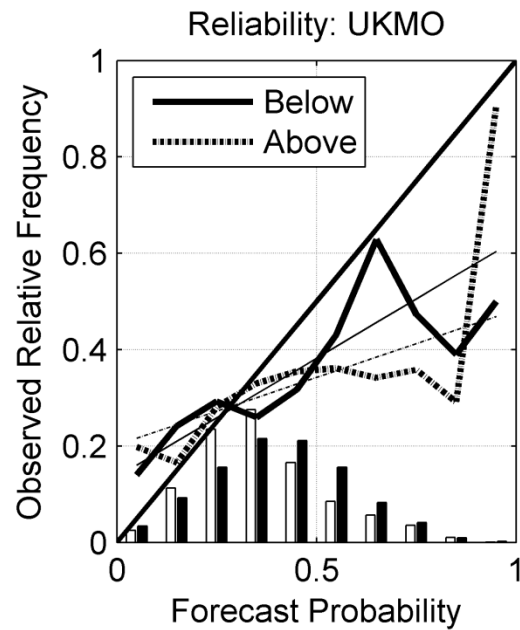
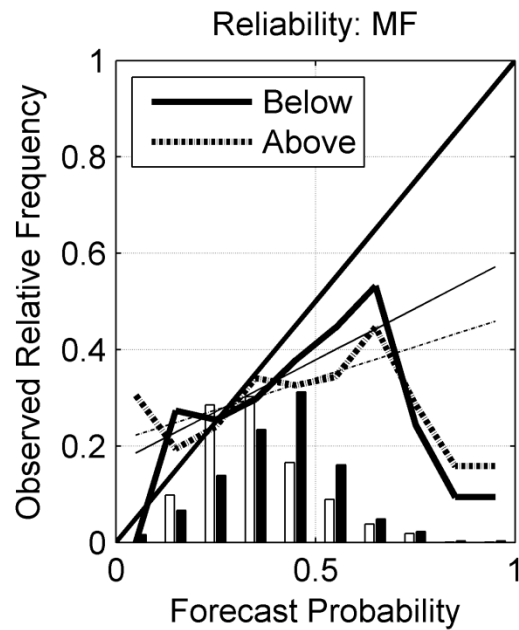
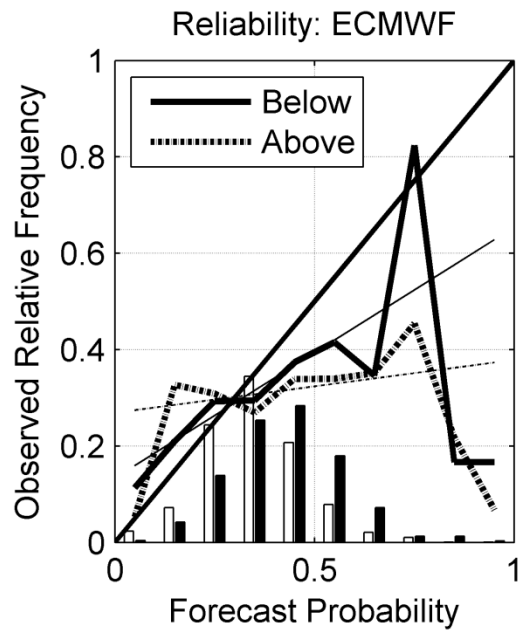
DJF Rainfall Area-Average (Africa south of 10°S)



COUPLED MODELS AND SEASONAL FORECASTING

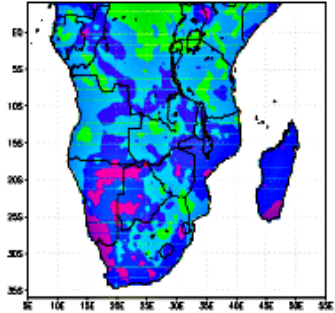
- ✘ ECHAM4.5
 - + Constructed analogue SST
 - + MOM3
 - + GML (slab)
- ✘ 850 hPa gp height fields downscaled to 94 districts
- ✘ Initial training period
 - + 1982/83 – 1995/96
- ✘ Retro-active test period
 - + 1996/97 – 2008/09



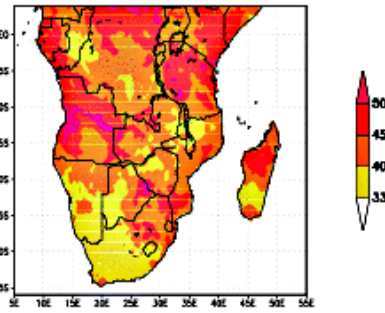


LATEST SEASONAL FORECASTS

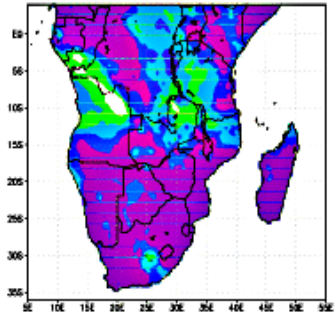
NOVEMBER–DECEMBER–JANUARY 2010/11
Above-Normal Rainfall



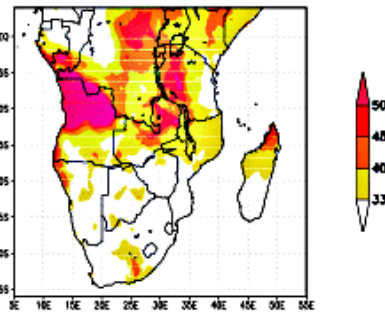
NOVEMBER–DECEMBER–JANUARY 2010/11
Below-Normal Rainfall



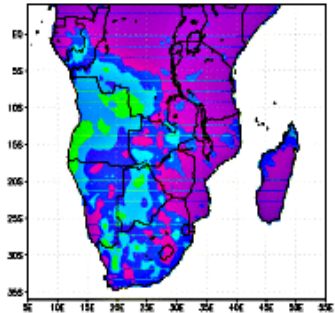
DECEMBER–JANUARY–FEBRUARY 2010/11
Above-Normal Rainfall



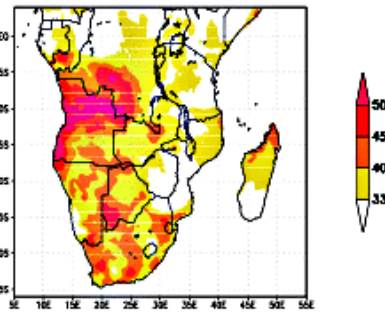
DECEMBER–JANUARY–FEBRUARY 2010/11
Below-Normal Rainfall



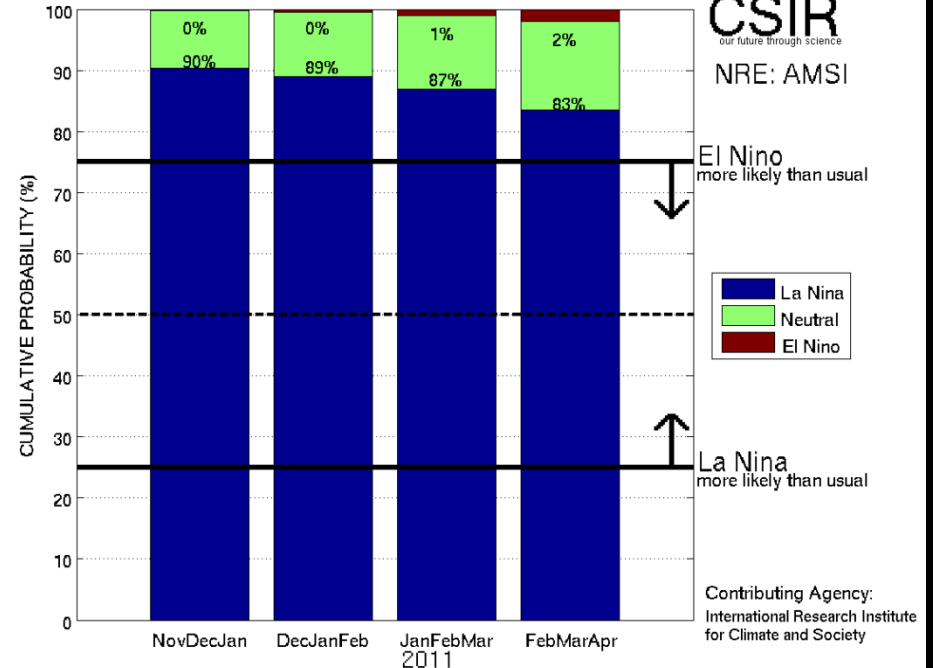
JANUARY–FEBRUARY–MARCH 2011
Above-Normal Rainfall



JANUARY–FEBRUARY–MARCH 2011
Below-Normal Rainfall

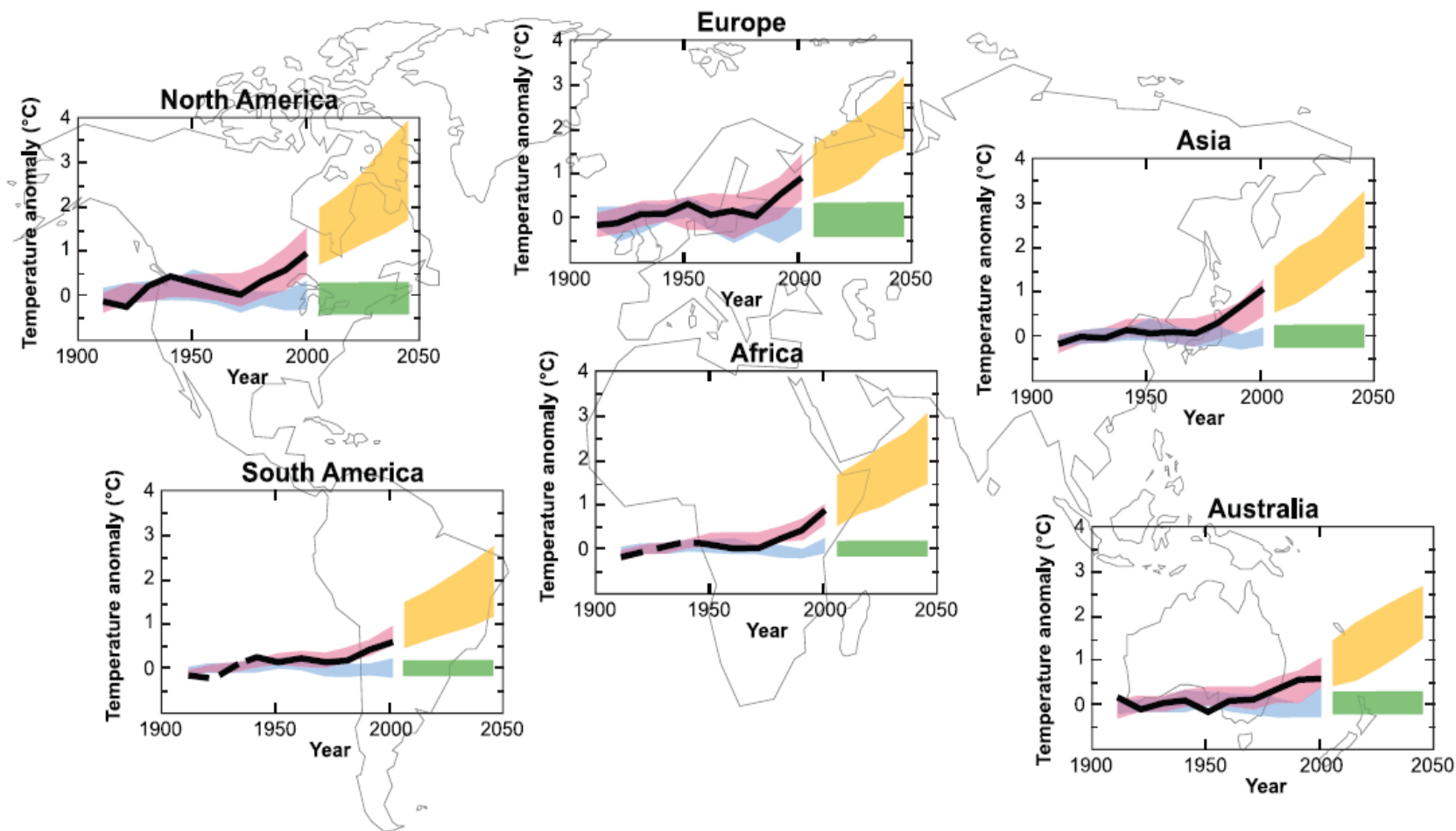


MULTI-MODEL PROBABILISTIC ENSO FORECASTS (NINO3.4 SST)
ISSUED ON: 08-Oct-2010



To find out how ENSO may affect the rainfall over southern Africa during the months ahead, please refer to the forecasts for SADC:
http://rava.qscons.net/themes/climate_template/

CLIMATE CHANGE PROJECTIONS



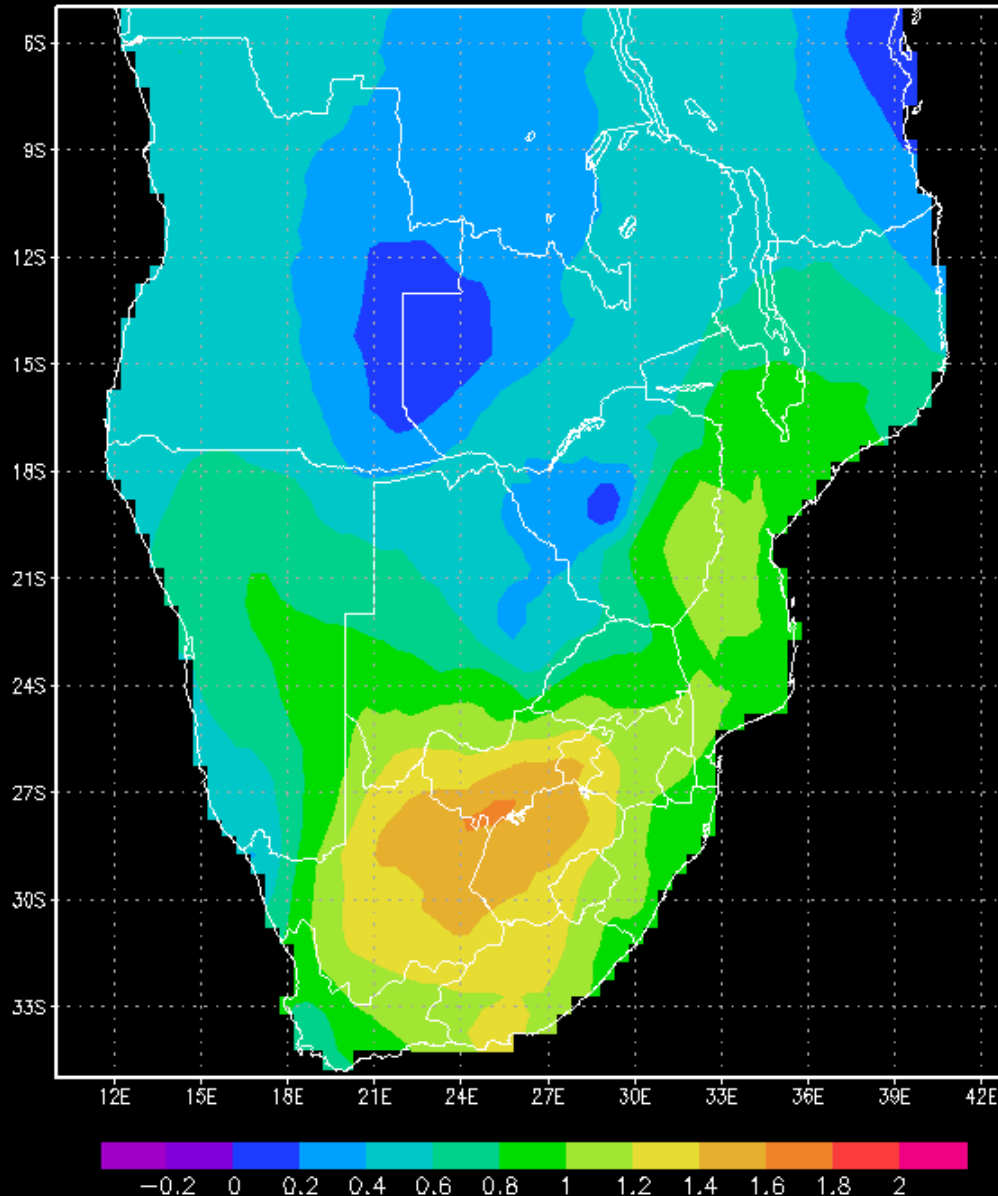
©IPCC 2007: WG1-AR4



Yearly temperature trends 1901-2002 (CRU)

* The largest trends occur over the central interior of South Africa – temps rising at more than twice the global average temp increase

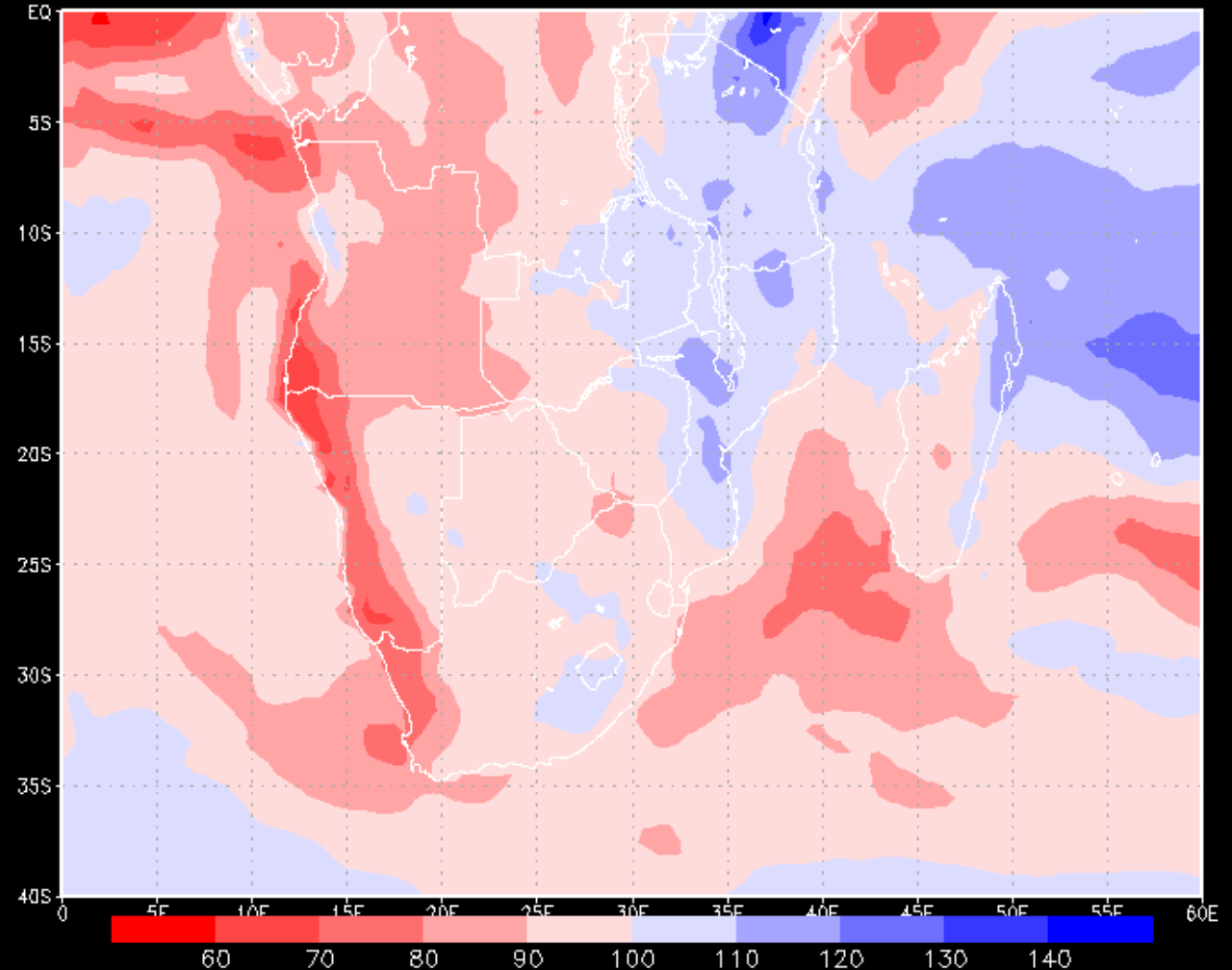
• Warming trend weaker over coastal areas



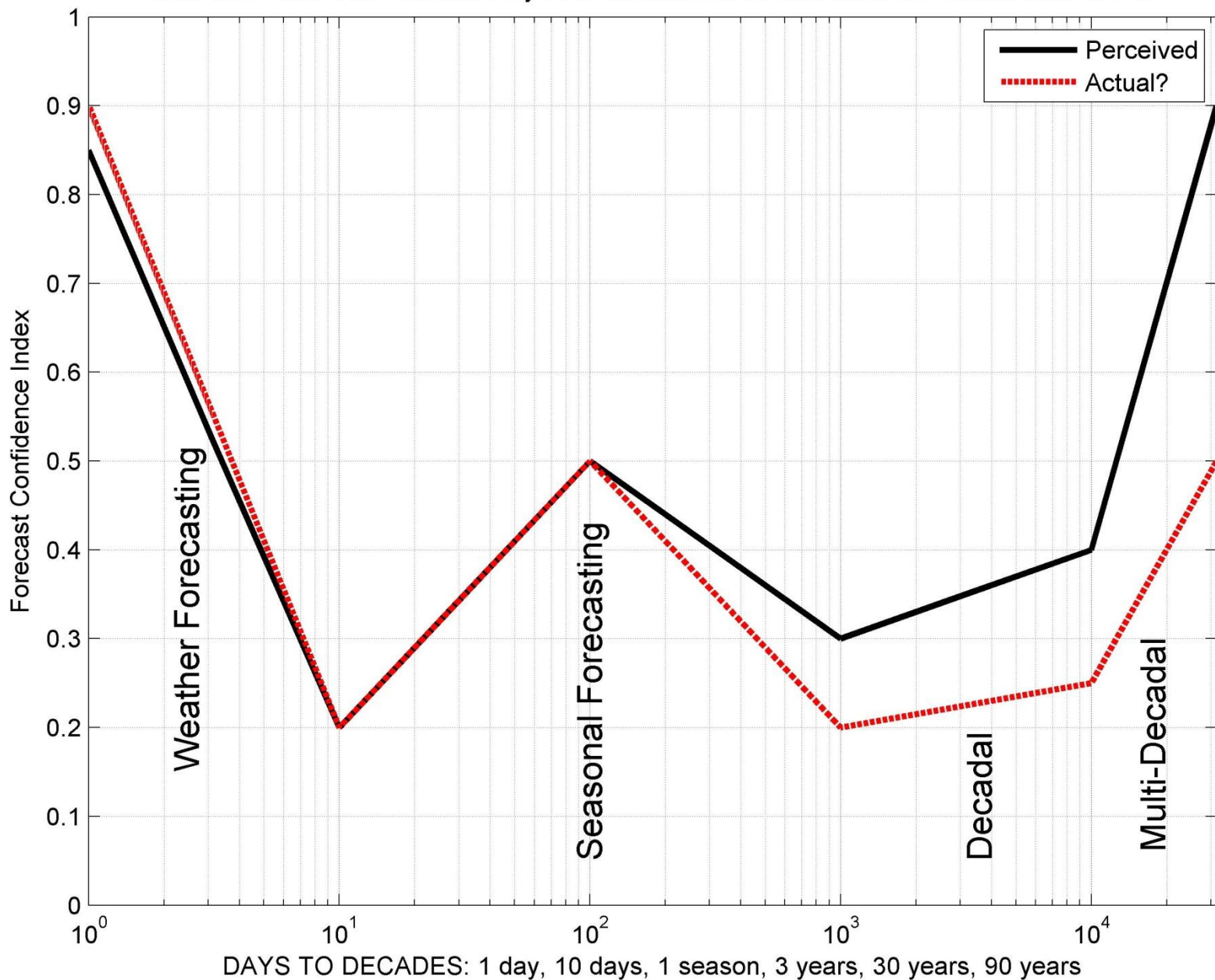
Change in yearly rainfall totals (%) over southern Africa

- A generally warmer and drier climate
- A shorter and more intense summer rainfall season
- Cloud bands displaced westwards in spring and autumn
- Drier winter rainfall region

C-CAM simulation of change in average yearly rainfall (%)



Real and Perceived Forecast/Projection Confidence for Mid-Summer Rainfall over SADC



Weather and Climate Modelling at the NRE

Climate Studies and Modelling (CSM)

Component 1

High-resolution
Weather
Forecasting

(1-10 days)

Component 2

Long-Range
Forecasting

(Weeks to months)

Component 3

Climate
Change
Projections

(Decadal and multi-decadal)

Component 4

Model
Development

Operational Forecasting

SEAMLESS FORECASTING: Using common forecast systems to predict for multiple time scales

FINAL THOUGHTS

- ✘ Forecasts for various time-ranges are subject to different forcings and have different skill levels
- ✘ Forecast models can provide useful forecasts
- ✘ Forecast model development has resulted in improved forecasts, also of extreme events
- ✘ The South African modelling community, in partnership with international modellers, is constantly improving on operational forecast systems